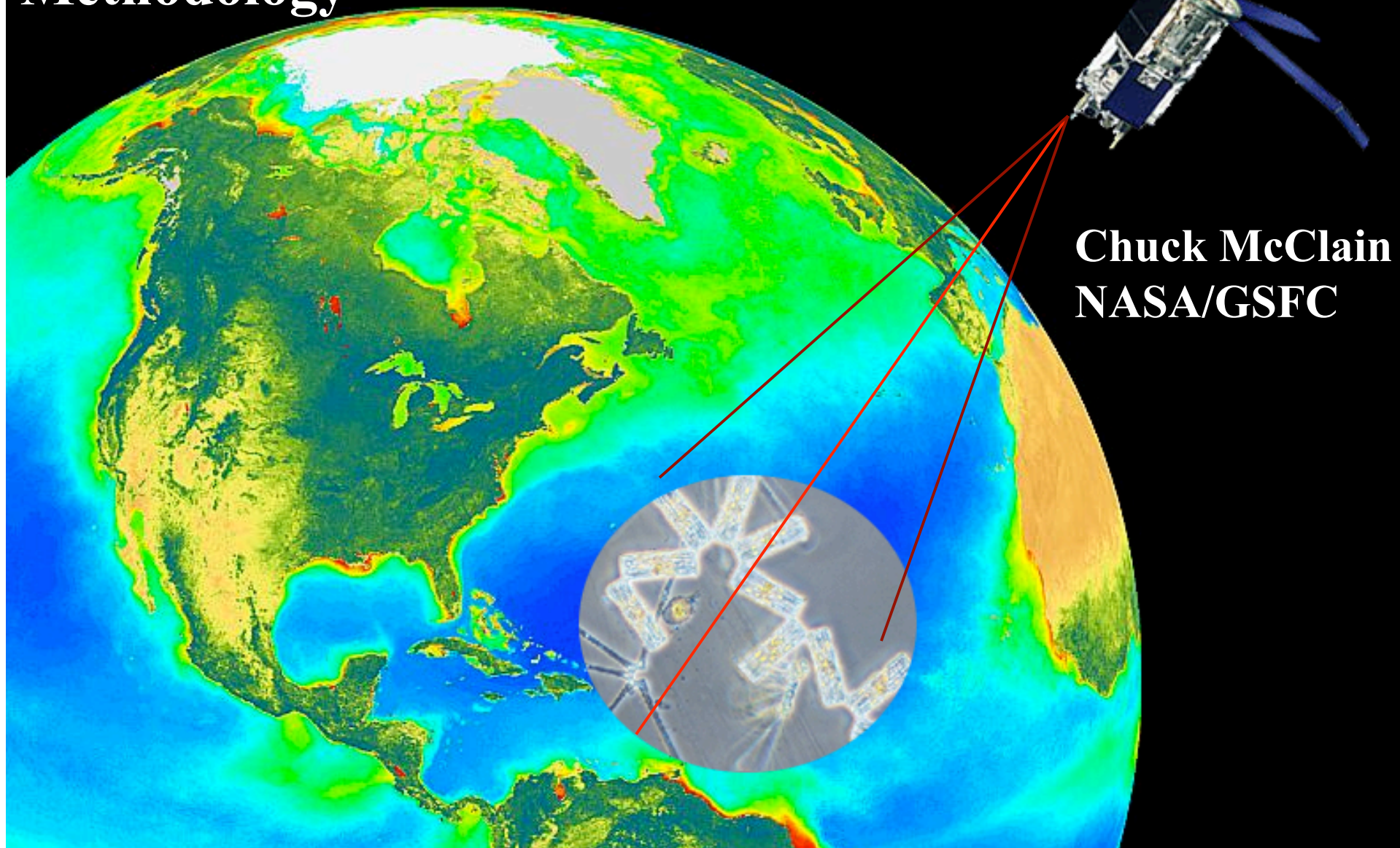


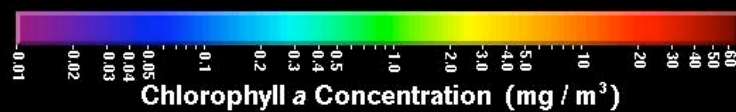
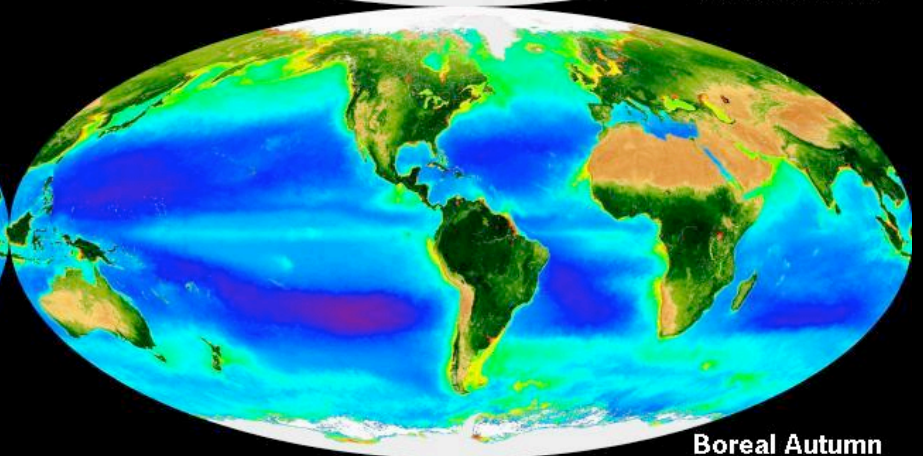
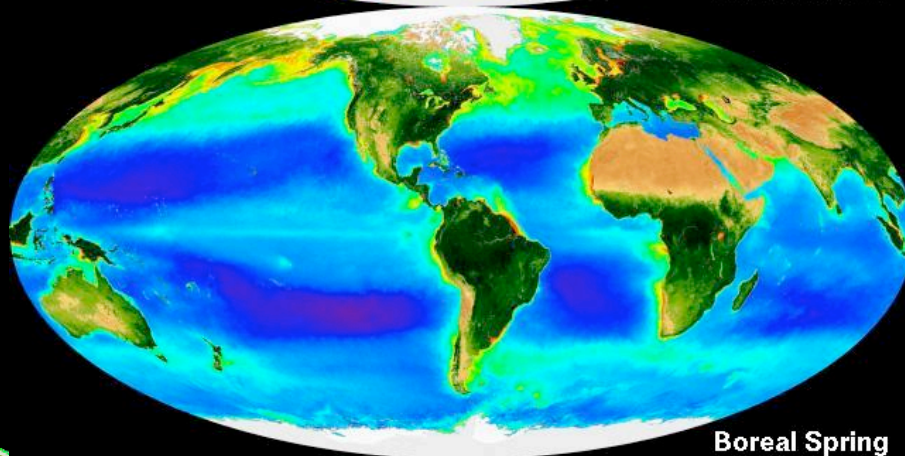
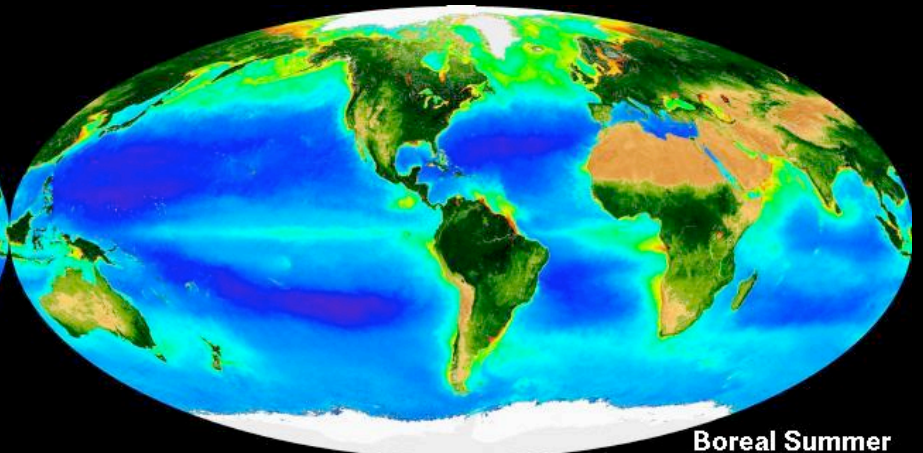
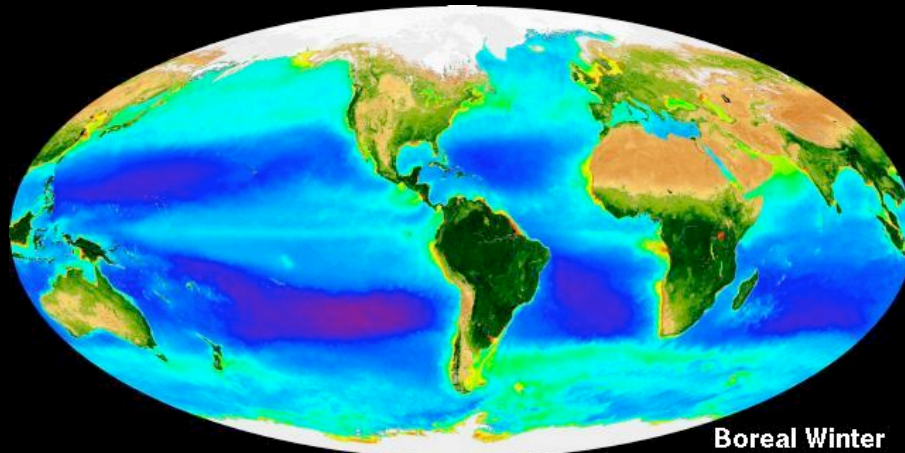
Review of the Ocean Biogeochemistry Calibration and Aerosol Correction Methodology



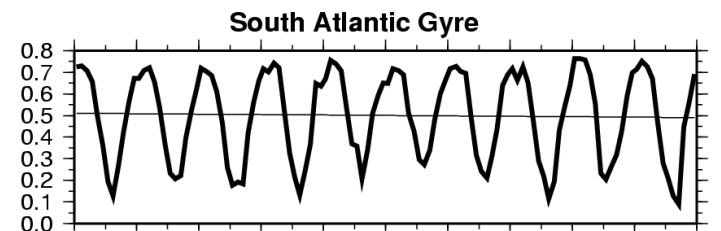
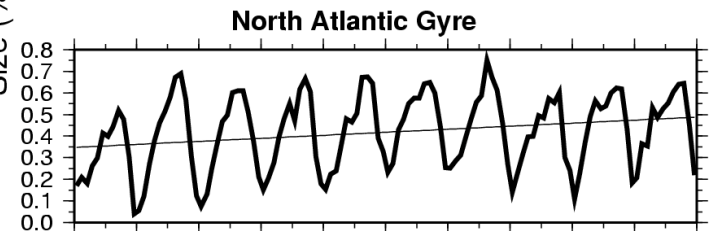
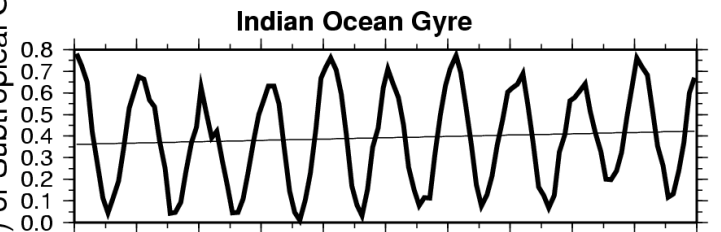
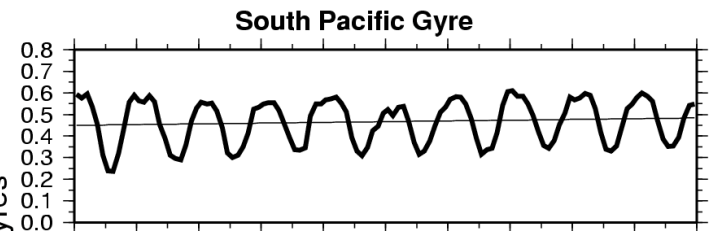
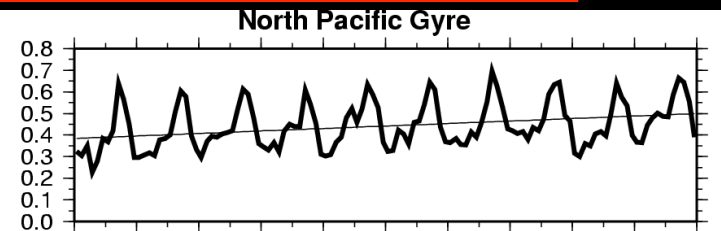
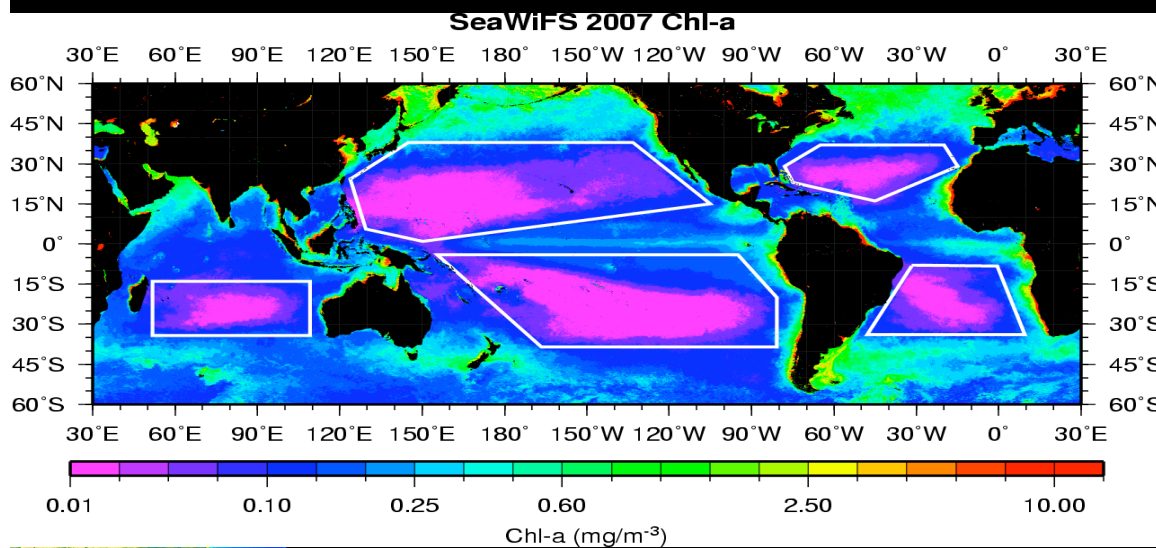
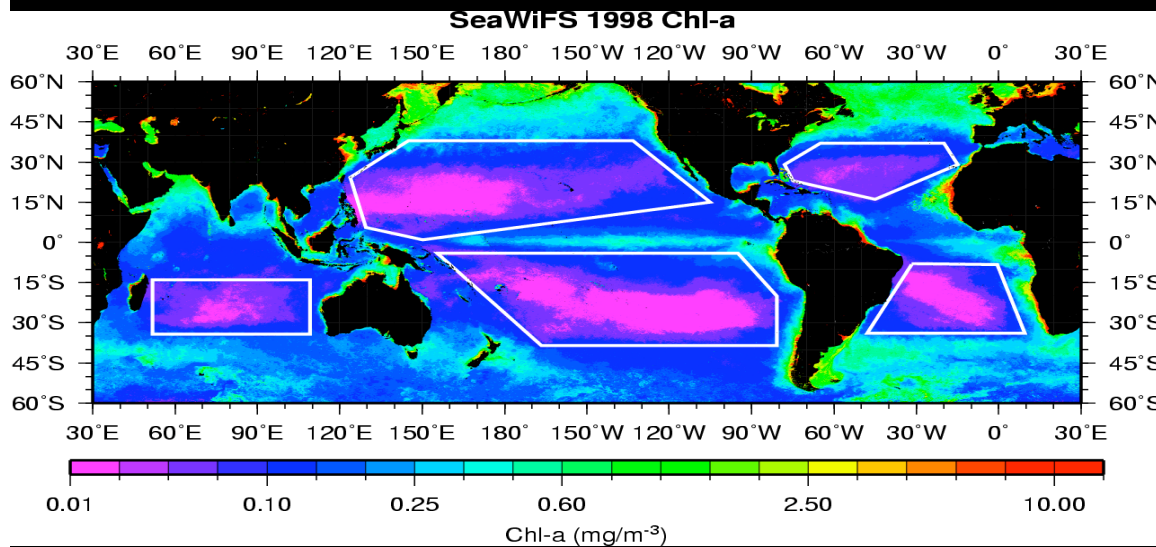
**Chuck McClain
NASA/GSFC**

Seasonal Global Biospheres

SeaWiFS Climatologies (1997-2002)



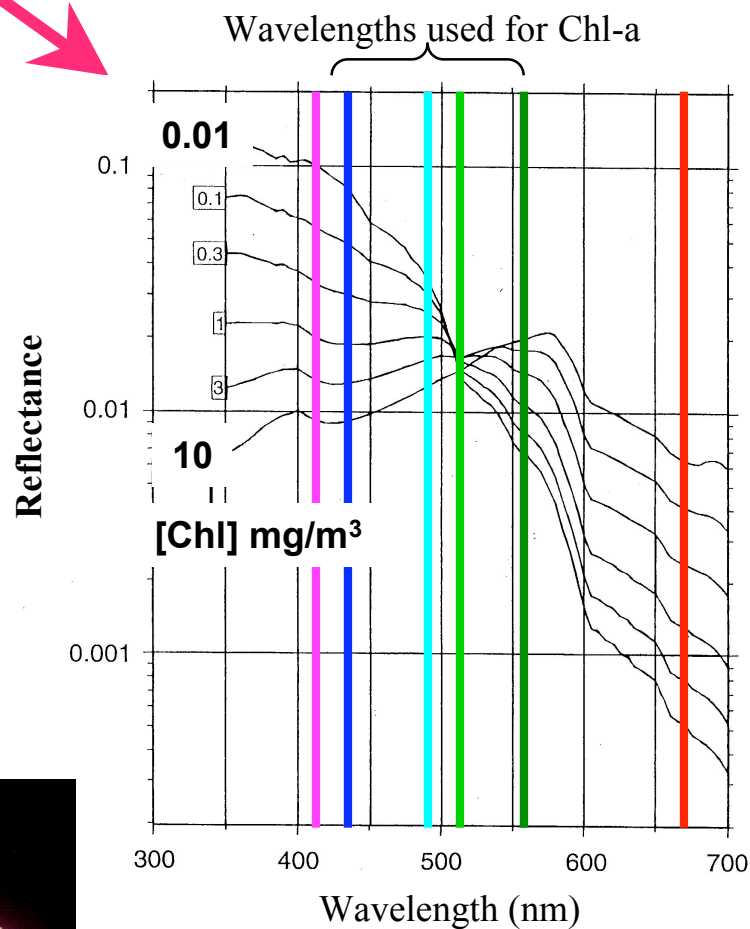
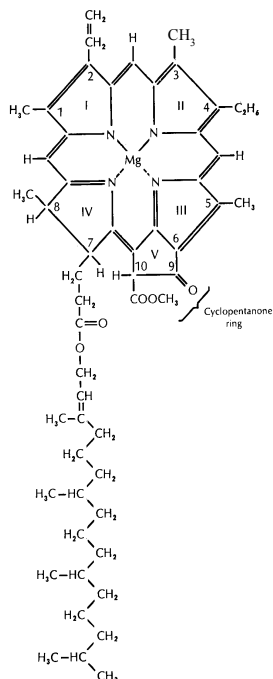
Decadal Scale Biological Variability



1998 1999 2000 2001 2002 2003 2004 2005 2006 2007 2008

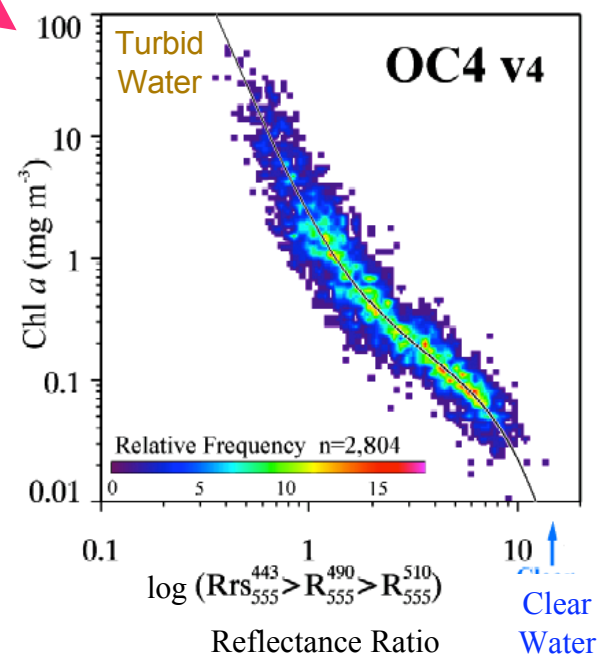
Quantifying Phytoplankton from Space: Chl-a Algorithm

Chlorophyll-a

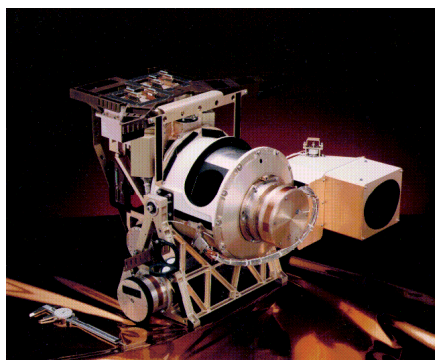


From chlorophyll absorption to
chlorophyll concentration via optics

*Chlorophyll Algorithm:
Statistical "Band-Ratio"
Regression*



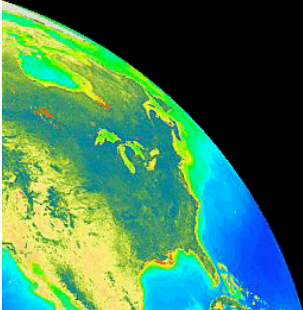
SeaWiFS



*Marine Spectral Reflectance
vs. Chlorophyll-a*

NASA Ocean Color Accuracy Goals

- **Sensor radiometric calibration**
 $\pm 0.5\%$ absolute
- **Water-leaving radiances**
 $\pm 5\%$ absolute
- **Chlorophyll-a**
 $\pm 35\%$ over range of 0.05-50.0 mg/m³



Note: NPOESS Climate Data Record accuracy goals are different.

SeaWiFS, MODIS, & VIIRS

- **SeaWiFS**

- Rotating telescope
- **412, 443, 490, 510, 555, 670, 765, 865 nm bands**
- 12 bit digitization truncated to 10 bits on spacecraft
- 4 focal planes, 4 detectors/band, 4 gain settings, bilinear gain configuration
- Polarization scrambler: sensitivity at 0.25% level
- Solar diffuser (daily observations)
- Monthly lunar views at 7° phase angle via pitch maneuvers

- **NPP/VIIRS (Ocean Color)**

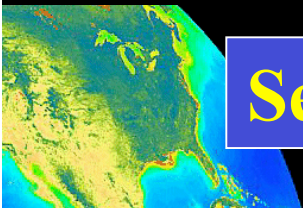
- SeaWiFS-like rotating telescope
- MODIS-like focal plane arrays (16 detectors/band)
- 12 bit digitization
- No polarization scrambler
- Solar diffuser with stability monitor
- 7 OC bands (**412, 445, 488, 555, 672, 746, 865 nm**)
 - Dual gains except 746 nm (single gain)
- Monthly lunar views at 55° phase angle via space view port with roll maneuvers (feasible, but not approved)

- **MODIS (Ocean Color)**

- Rotating mirror
- **413, 443, 488, 531, 547, 667, 678, 748, 870 nm bands**
 - Single gain (NIR saturation)
- 12 bit digitization
- 4 focal planes (7-11 bands each)
 - OC Visible: 412-547 nm (5 bands-10 detectors each)
 - OC NIR: 667-869 (4 bands-10 detectors each)
- No polarization scrambler: sensitivity at ~3% level
- Spectral Radiometric Calibration Assembly (SRCA)
- Solar diffuser (observations every orbit), Solar Diffuser Stability Monitor (SDSM)
- Monthly lunar views at 55° phase angle via space view port

“Nominal” common bands

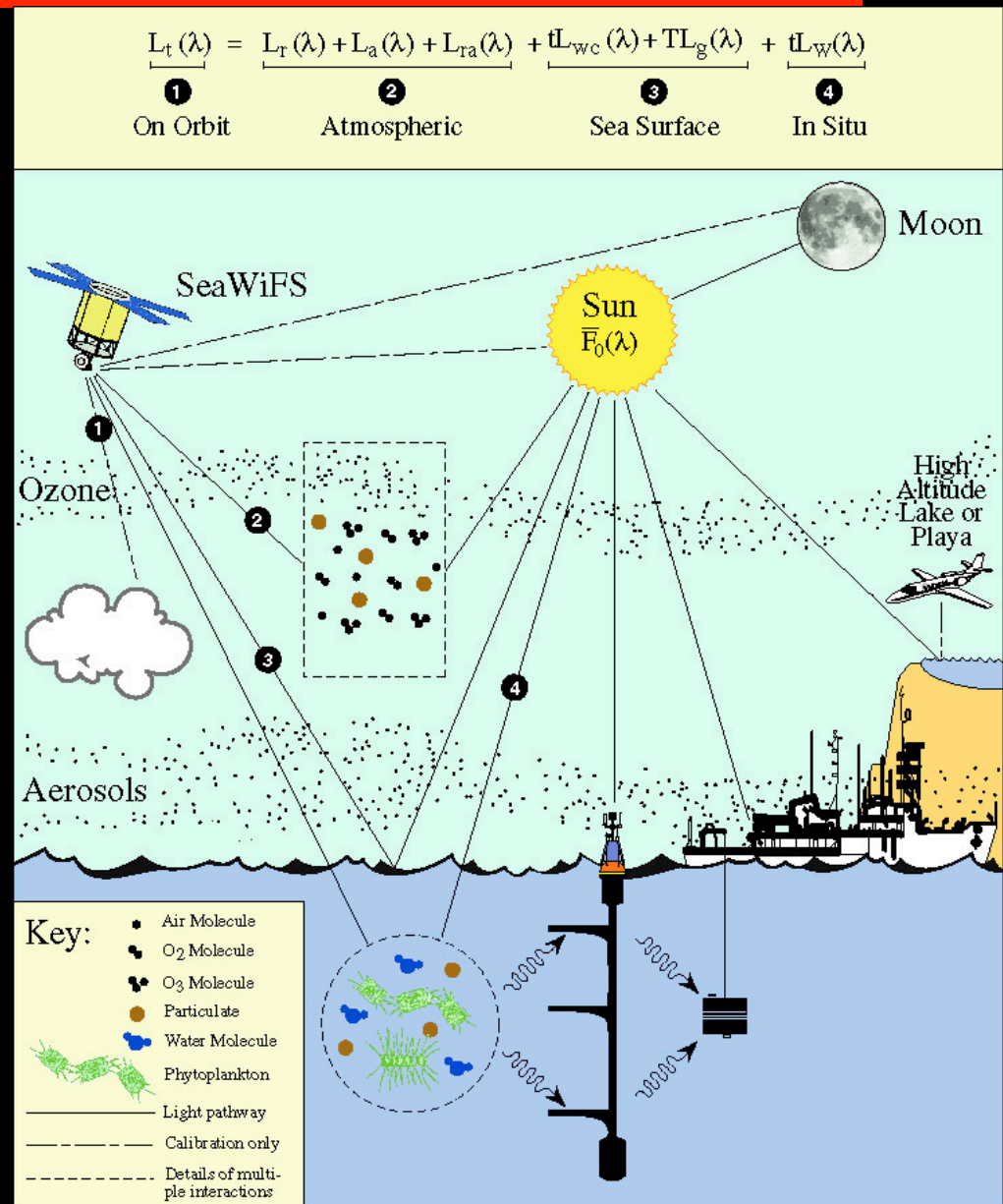
Sensor designs & performance are never identical.



Calibration/Validation Paradigm

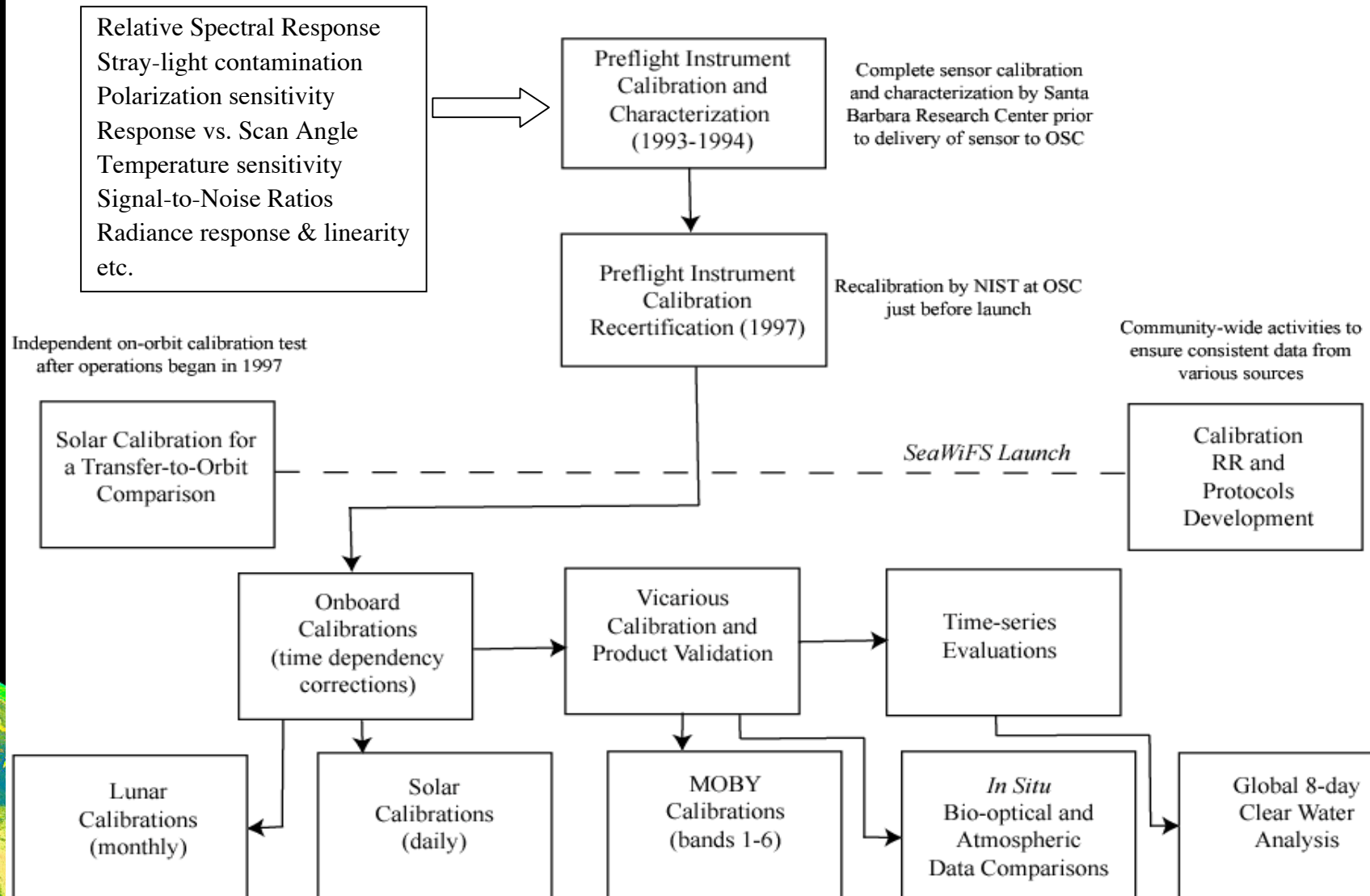
Program Elements:

- **Laboratory** - prelaunch sensor calibration & characterization
- **On-orbit** - solar and lunar observations are used to track changes in sensor response
- **Field** - comparison of satellite data retrievals to in-water, above-water and atmospheric observations
 - Vicarious calibration - adjust instrument gains to match water-leaving radiances
 - Product validation (water-leaving radiances, chl-a, etc.)

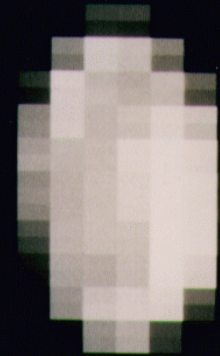


Ocean Color Sensor Calibration Strategy

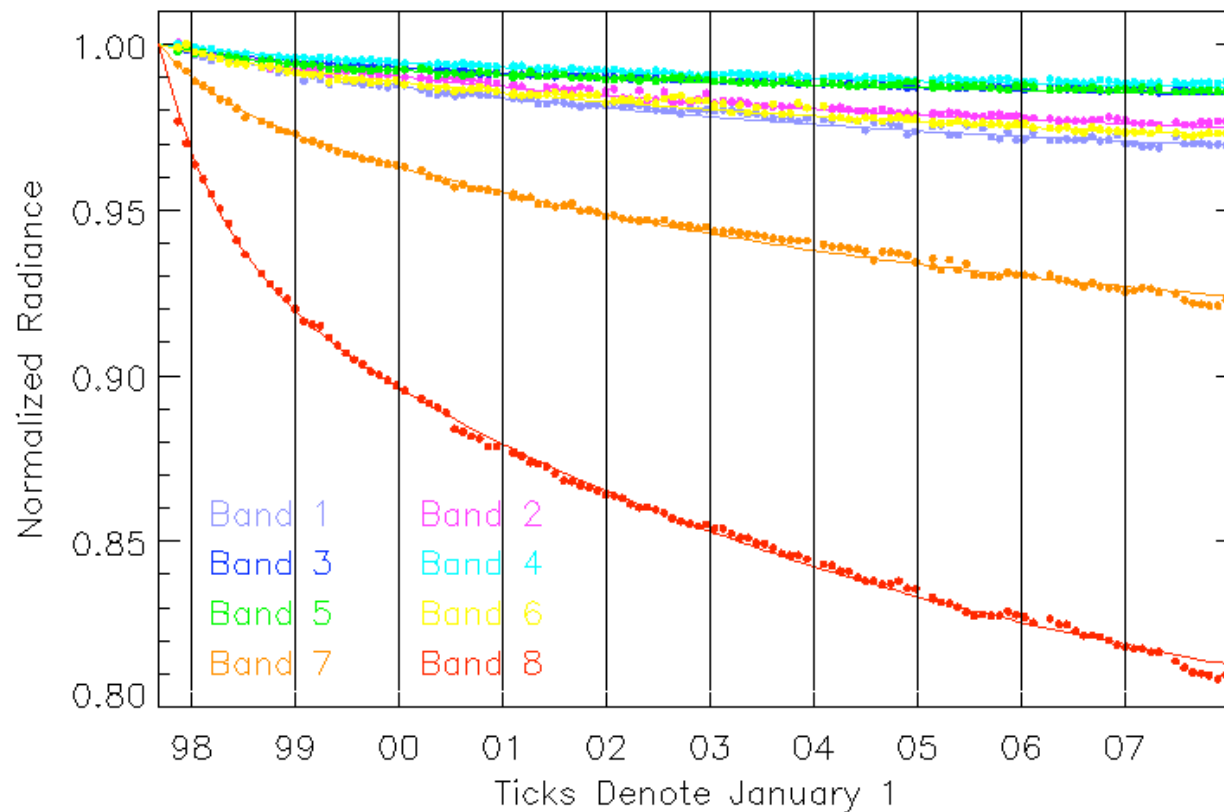
SeaWiFS Calibration Strategy



SeaWiFS Temporal Calibration

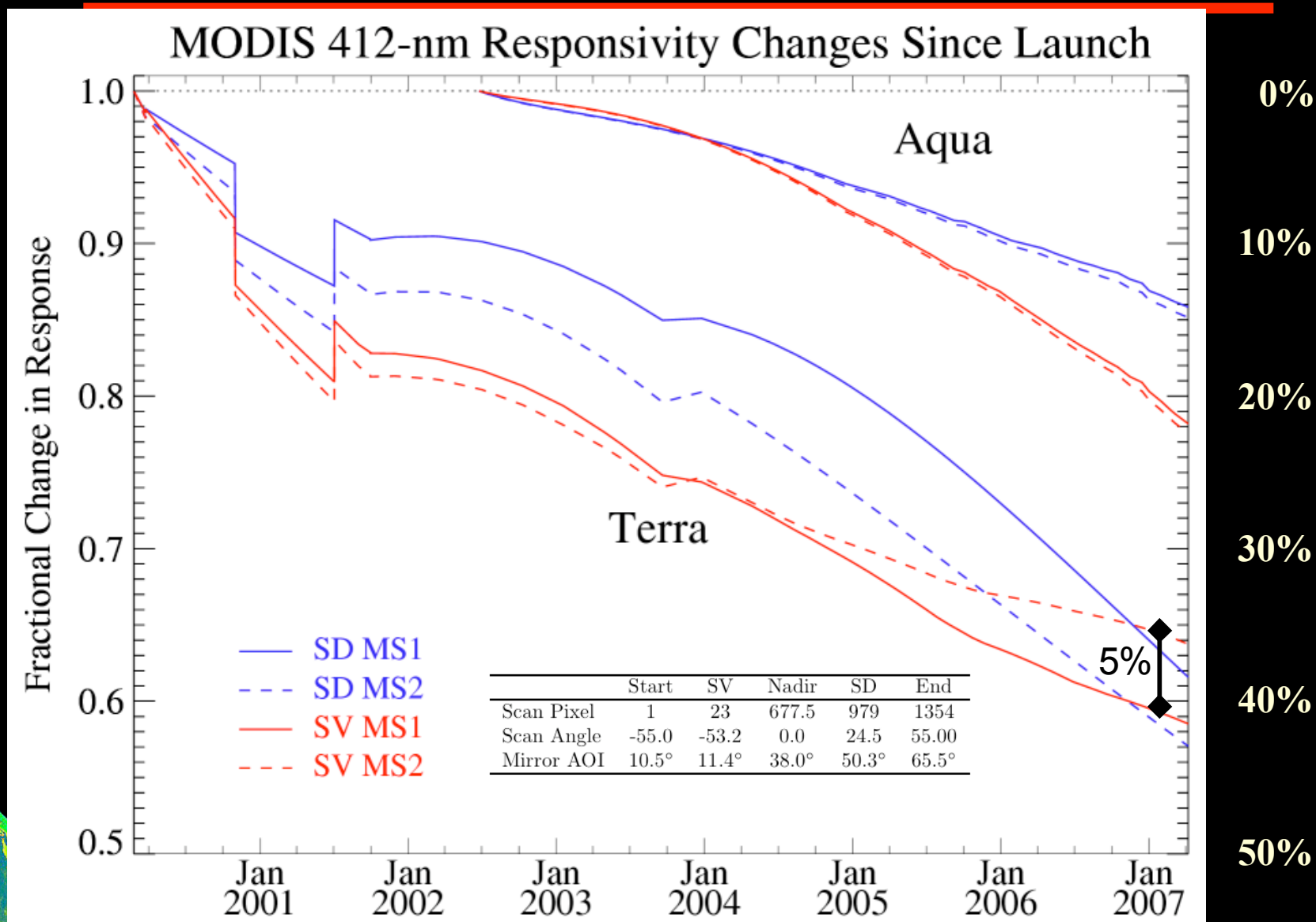


SeaWiFS Lunar Calibrations



SeaWiFS Band	SeaWiFS λ (nm)
1	412
2	443
3	490
4	510
5	555
6	670
7	765
8	865

MODIS Temporal Degradation at 412 nm Lunar and Solar Calibration Trends



OC Atmospheric Correction Basics

- Methodology from Gordon & Wang (1994)
 - Assumes NIR reflectance is negligible
 - Models Rayleigh multiple scattering (plane-parallel model)
 - Employs suite of 12 aerosol models (next slide)
 - Uses ratio of NIR aerosol radiances to determine aerosol model
 - Aerosol model extrapolation to visible bands to estimate aerosol radiance
- Glint and white cap radiances modeled using NCEP surface winds
- Gas absorption: Ozone, NO₂ (implemented, not operational), O₂ (SeaWiFS 765 nm only)



OC Atmospheric Correction Basics

$$L_t = [L_r + L_a + T_v L_g + t_{dv} (L_f + L_w)] t_{gv} t_{gs} f_p$$

L_t : measured top of atmosphere radiance

L_r : Rayleigh radiance

All terms λ -dependent

L_a : aerosol radiance

T_v : direct transmittance (sensor view)

L_g : glint radiance

t_{dv} : diffuse transmittance (sensor view)

L_f : foam radiance

L_w : water-leaving radiance

t_{gv} : absorbing gas transmittance (ozone; sensor view path)

t_{gs} : absorbing gas transmittance (ozone; solar path)

f_p : sensor polarization sensitivity correction factor

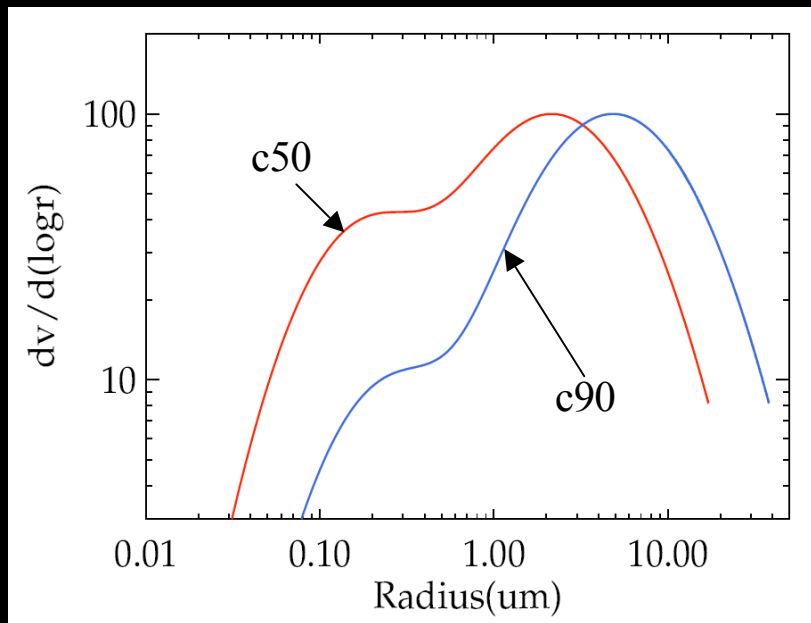
Aerosol Models for Atmospheric Correction

- Gordon & Wang's aerosol models based on Shettle and Fenn's (1979) models for tropospheric and oceanic aerosols
 - Twelve (12) aerosol models are used in operational processing (#'s refer to relative humidity)
 - Oceanic
O99
 - Maritime (1% oceanic and 99% tropospheric)
M99, M90, M70, M50
 - Coastal (0.5% oceanic and 99.5% tropospheric)
C99, C90, C70, C50
 - Tropospheric
T99, T90, T50

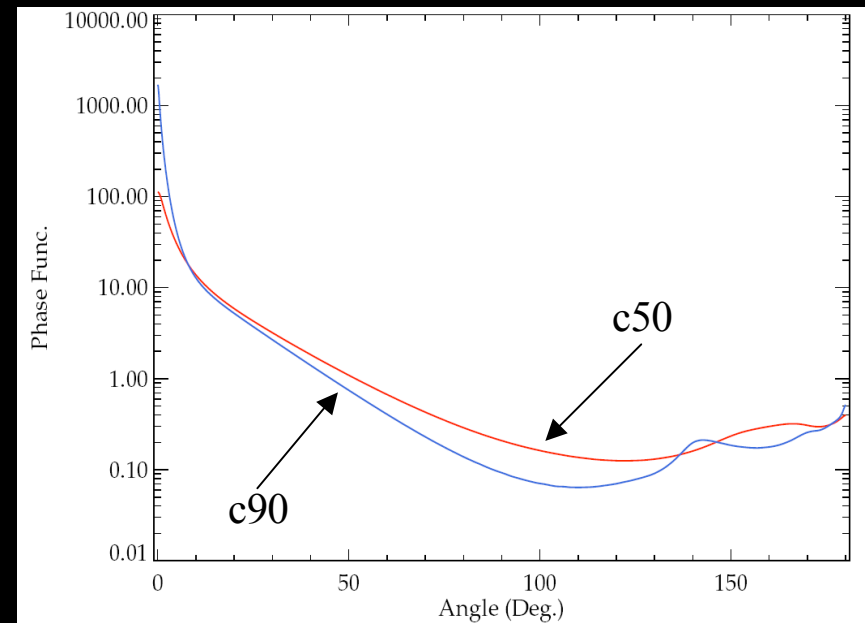


Some Properties of G&W Aerosol Models

Size Distribution



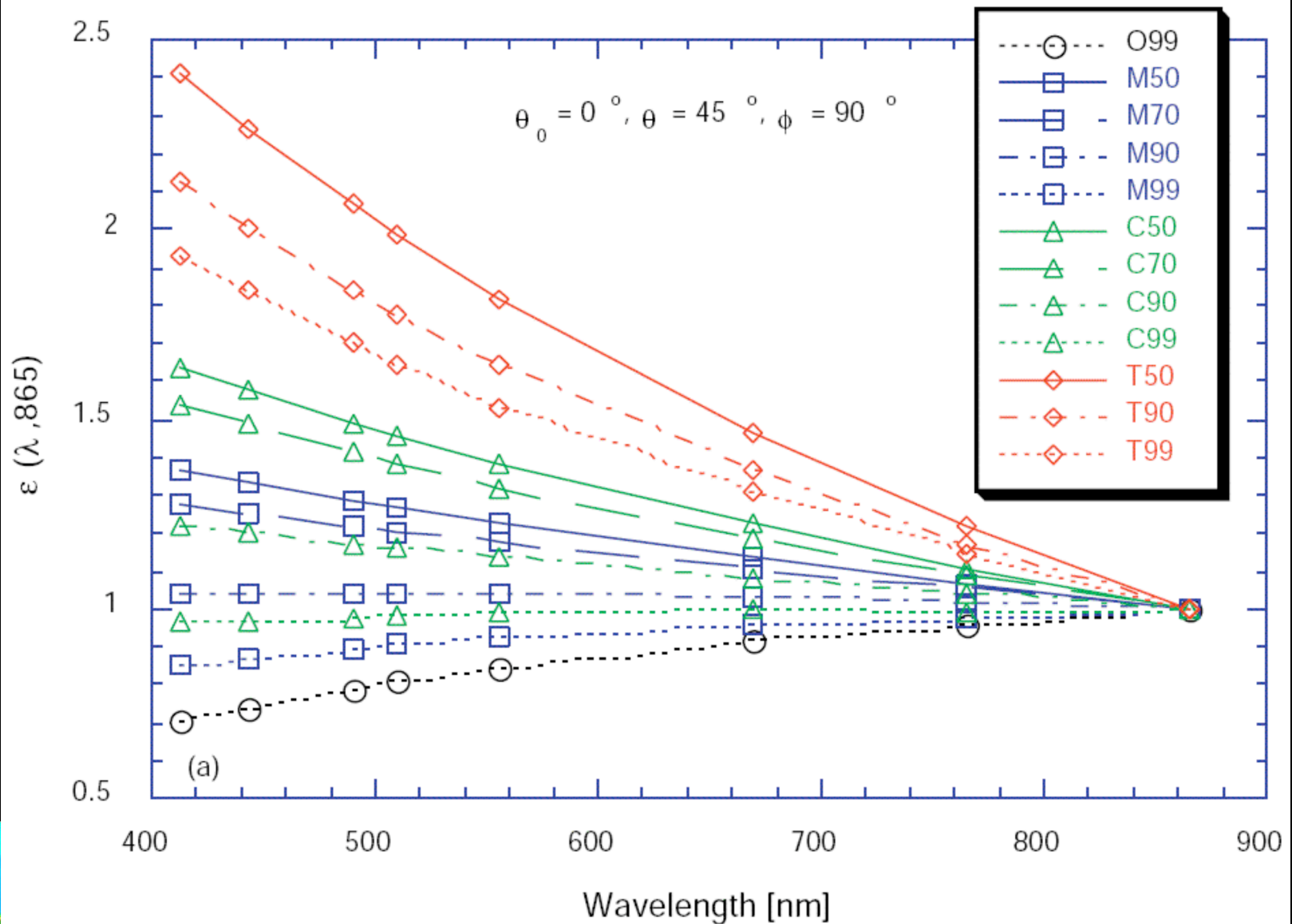
Phase Function



- Effective radius varies from 0.14 to 4.74 μm
- Single Scattering Albedo (SSA) varies from 0.93 (T50) to 1.0 (O99)

Epsilons for 12 Models

Solar Zenith = 0°, Sensor Zenith = 45°, Sensor Azimuth = 90°

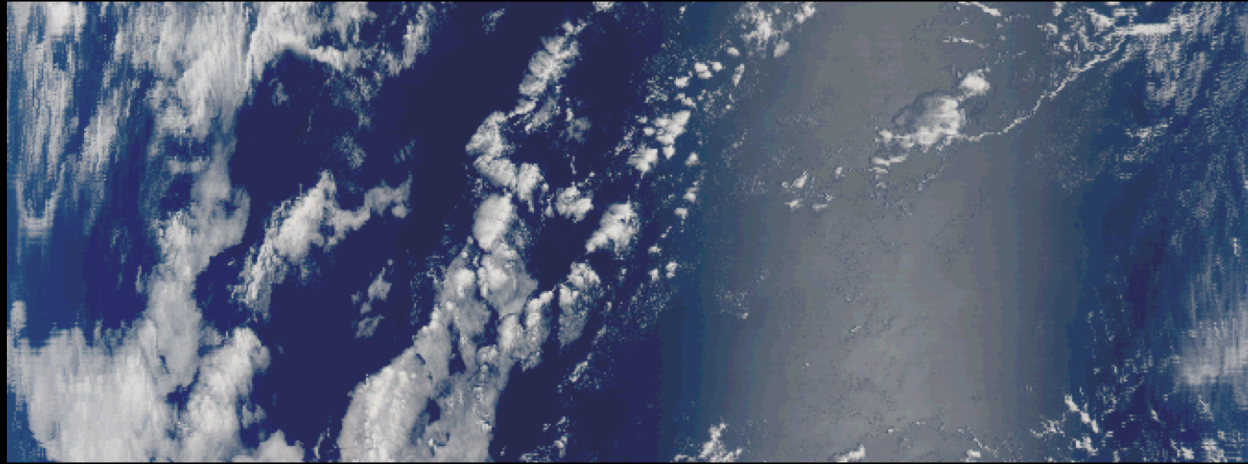


Computations of $\rho_{\lambda, 865}$ and $\rho_{\lambda, atm}$ Reflectances

- Convert the observed 765- and 865-nm reflectances into single scattering reflectances and compute $\rho_{765, 865} (obs, ss)$.
- Convert LUT multiple-scattering reflectances into single scattering reflectances using multiple to single scattering coeffs. and compute $\rho_{765, 865} (cal, ss)$ for each aerosol model.
- Select two models that bracket the observe $\rho_{765, 865}$. Also, compute $\rho_{\lambda, 865} (cal, ss)$ for both models.
- Since $\rho_{\lambda, 865} (cal, ss) = [\rho_{\lambda} (cal, ss) / \rho_{865} (cal, ss)]_{LUT} = [\rho_{\lambda} (obs, ss) / \rho_{865} (obs, ss)]$, compute $\rho_{\lambda} (obs, ss)$
- Convert $\rho_{\lambda} (obs, ss)$ into $\rho_{\lambda, obs} (obs, ms)$ using single to multiple scattering coeff.
- Combine $\rho_{\lambda} (obs, ms)$ from the two distributions to get the best estimate of $\rho_{\lambda, atm}$

Surface Effects

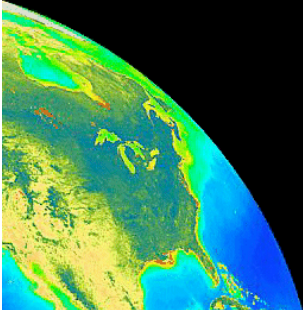
Sun Glint



White Caps



Corrections based
on statistical models
(wind & geometry)

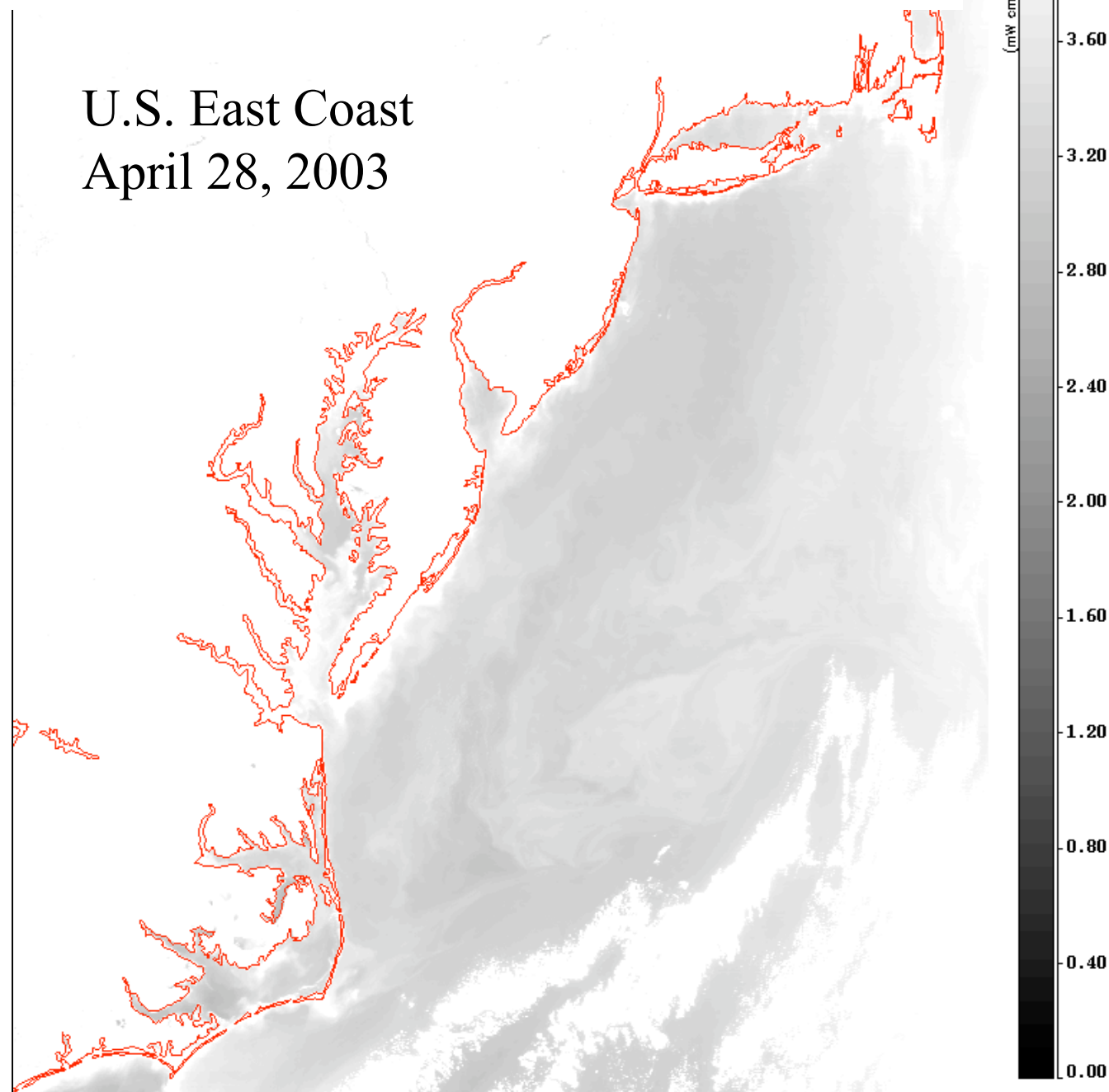


Atmospheric Correction: An Example

Green wavelength
551 nm

Total top-of-the-
atmosphere radiance

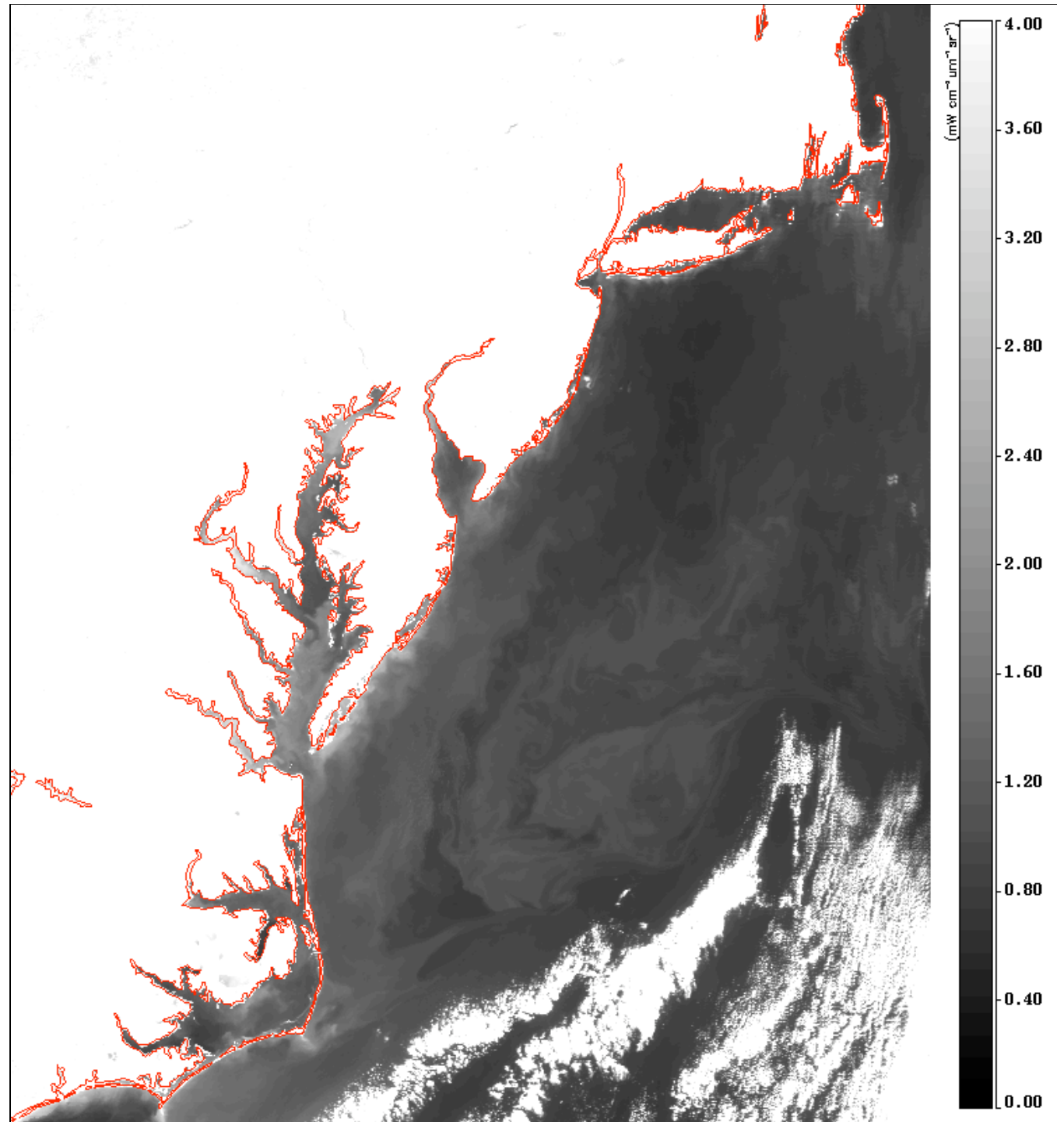
$0 - 4 \text{ mW/cm}^2 \text{ sr } \mu\text{m}$



Green wavelength
551 nm

Total top-of-the-
atmosphere radiance
corrected for ozone
absorption and
Rayleigh (gas
molecule) scattering

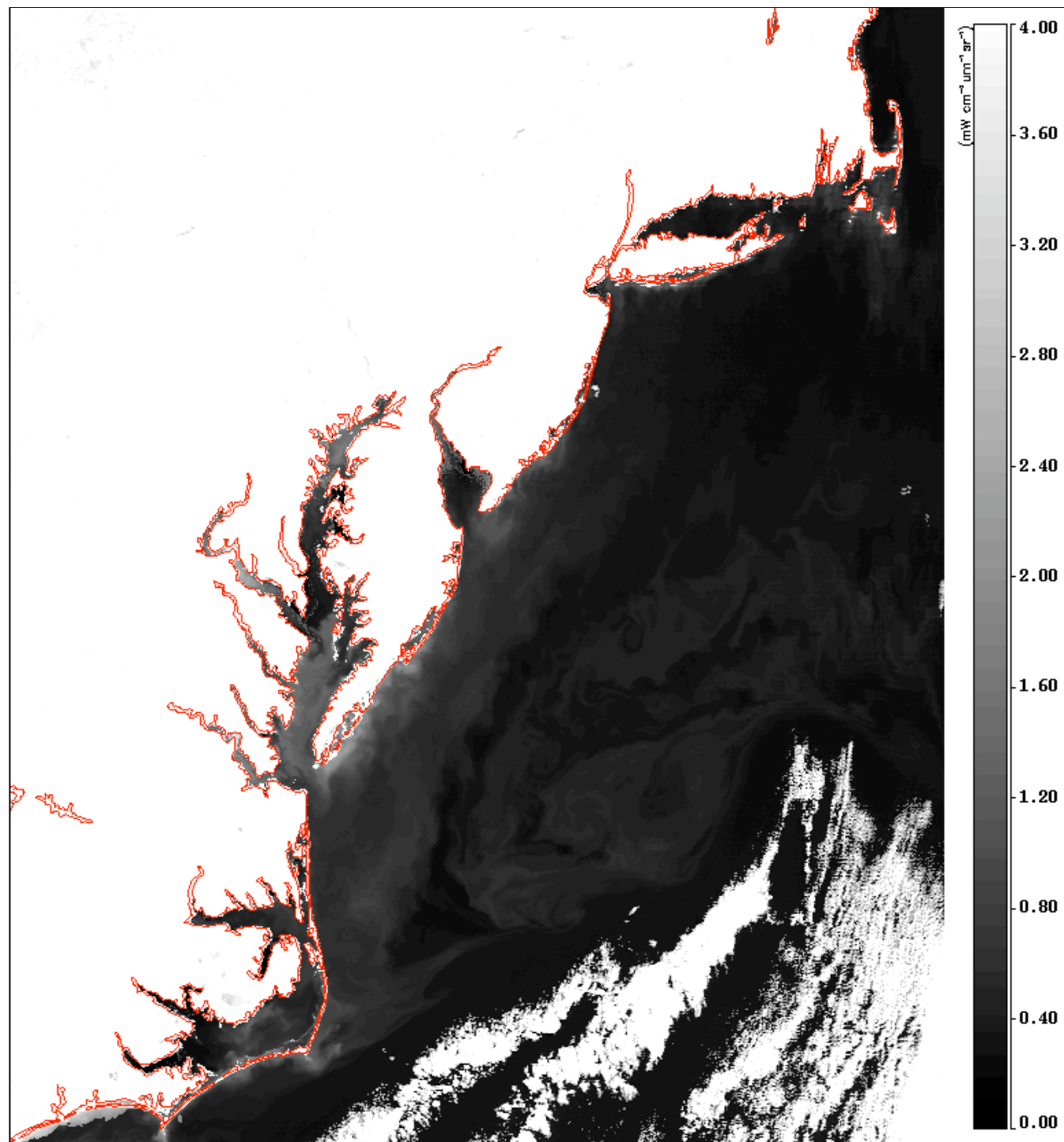
$0 - 4 \text{ mW/cm}^2 \text{ sr } \mu\text{m}$



Green wavelength
551 nm

Total top-of-the-
atmosphere radiance
corrected for ozone
absorption, Rayleigh
& aerosol scattering

0 – 4
 $\frac{\text{mW}}{\text{cm}^2 \text{ sr } \mu\text{m}}$

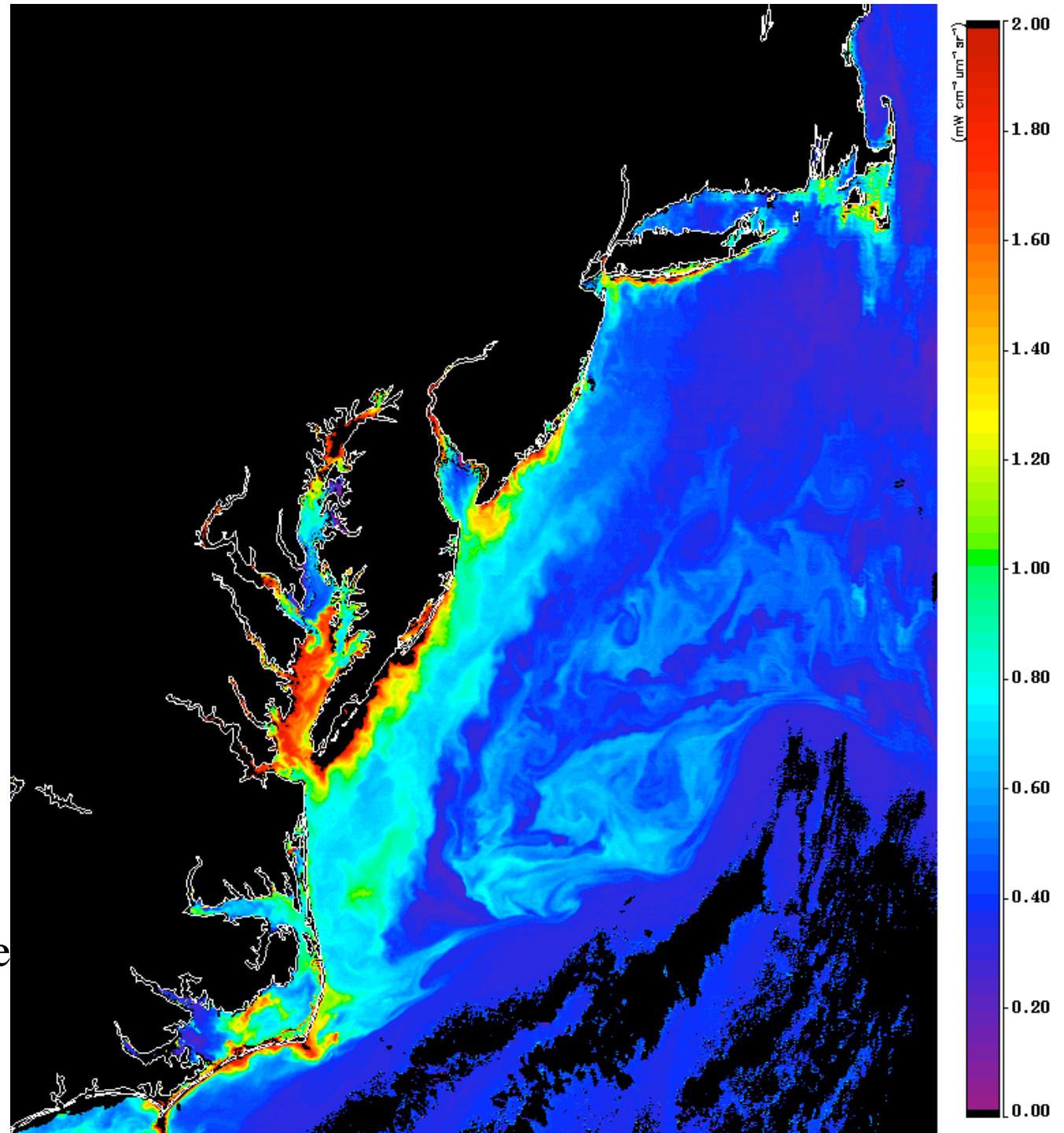


Green wavelength
551 nm

Normalized water-
leaving radiance

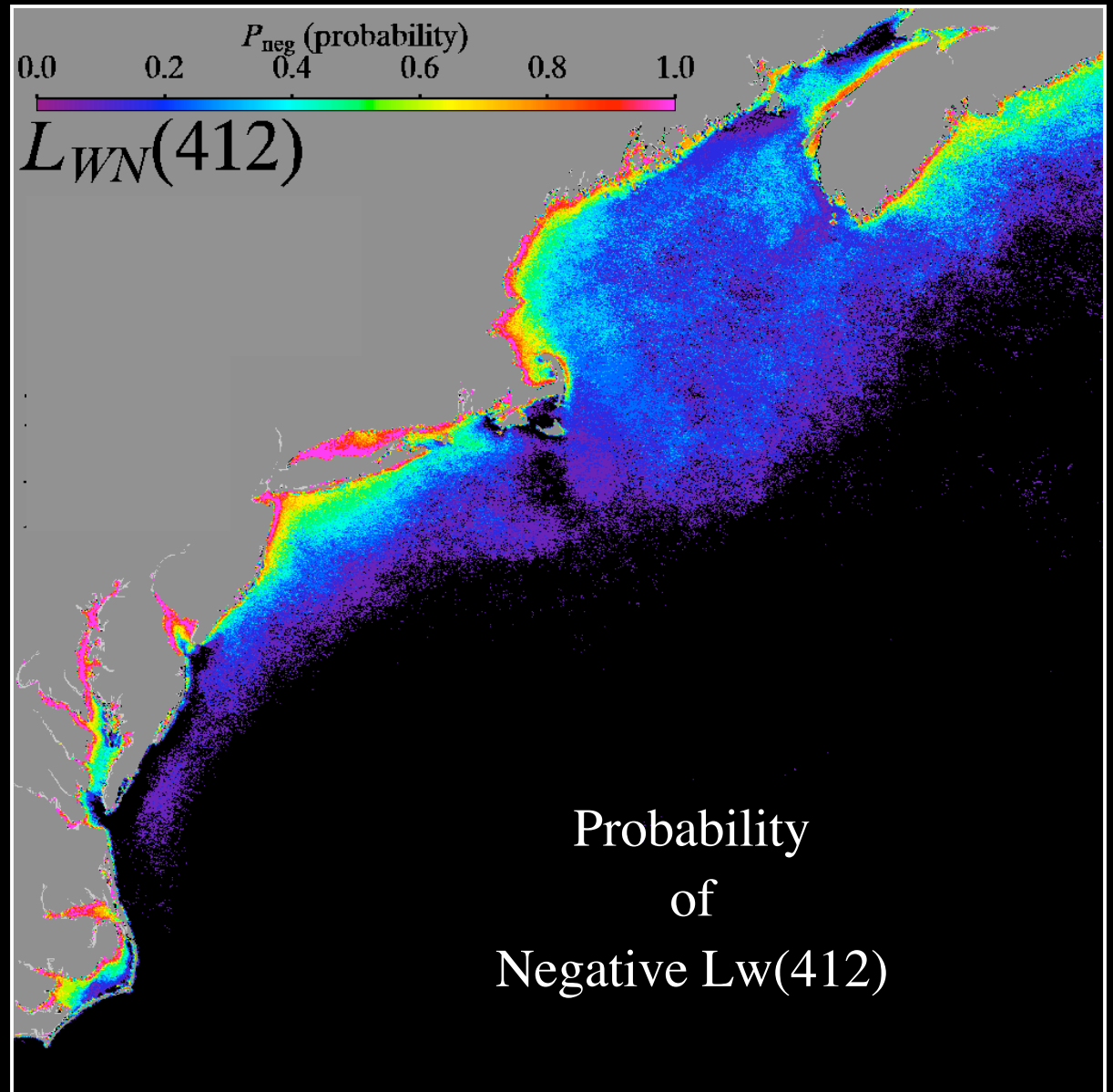
$$0 - 2 \text{ mW/cm}^2 \text{ sr } \mu\text{m}$$

Normalized: Lw is
transformed to radiance
normal to the surface
taking into account the
ocean BRDF



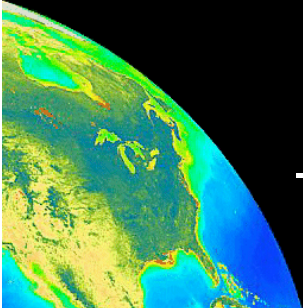
Atmospheric Correction: Negative Lw's

Certain coastal regions (e.g., NE U.S. & Southern California) are problematic due to aerosol types not represented in model suite & lack of NO₂ correction.



NIR Calibration Strategy

- Assume 865 & 869 (SeaWiFS, MODIS) nm calibrations after temporal degradation corrections are accurate
 - Visible band vicarious calibration insensitive to NIR calibration errors $< 5\%$
- Assume a fixed aerosol type/model for South Pacific and South Indian Ocean sites and adjust 765 & 748 nm band gains to yield “correct” aerosol radiance
 - The “M90” model (Gordon & Wang, 1994)
 - Properties similar to those inferred from the Tahiti AERONET site
 - Gain adjustment magnitude $\sim 2\text{-}3\%$ for SeaWiFS

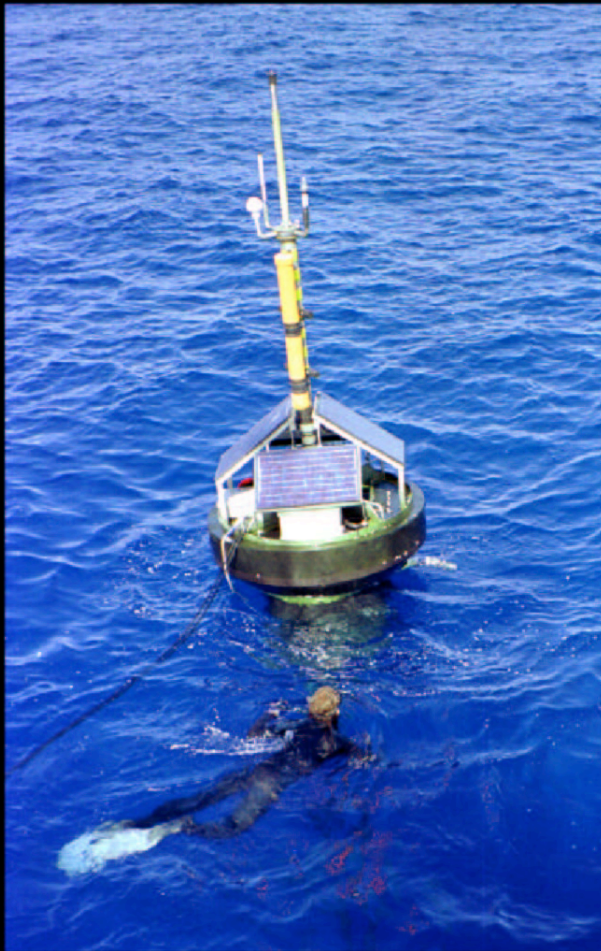


Vicarious Calibration: Visible

- Match satellite observations with surface L_w 's
- Apply atmospheric correction scheme
 - Determine aerosol properties
- Propagate L_w 's to top of atmosphere to estimate a “correct” L_t
- Compute calibration gain factor as ratio of “correct” L_t / measured L_t
- Average individual gain factors to produce mission-average gain factor

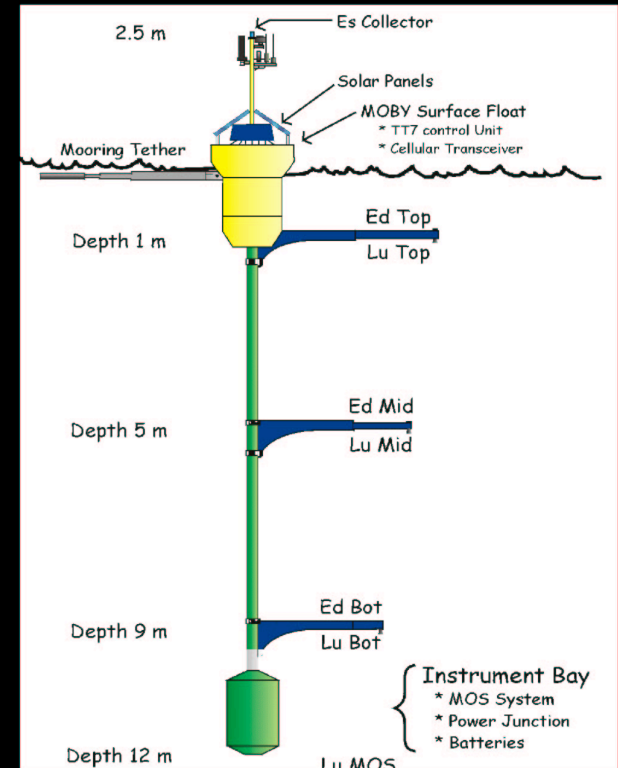


Vicarious Calibration: Visible

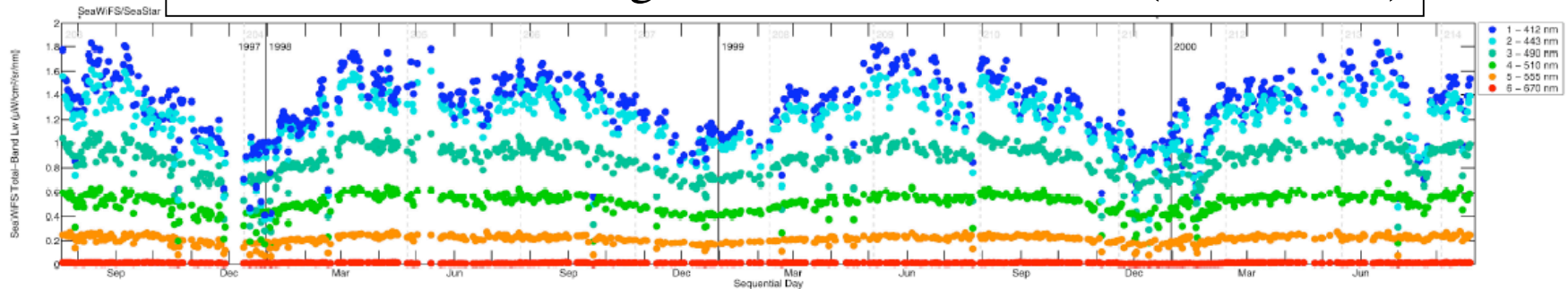


Marine Optical Buoy:

- Deep, clear water site (Lanai, Hawaii)
- Marine atmosphere
- Relatively cloud free
- Low latitude: year-round satellite coverage at useful viewing geometries
- Glint contamination – nontilting sensors

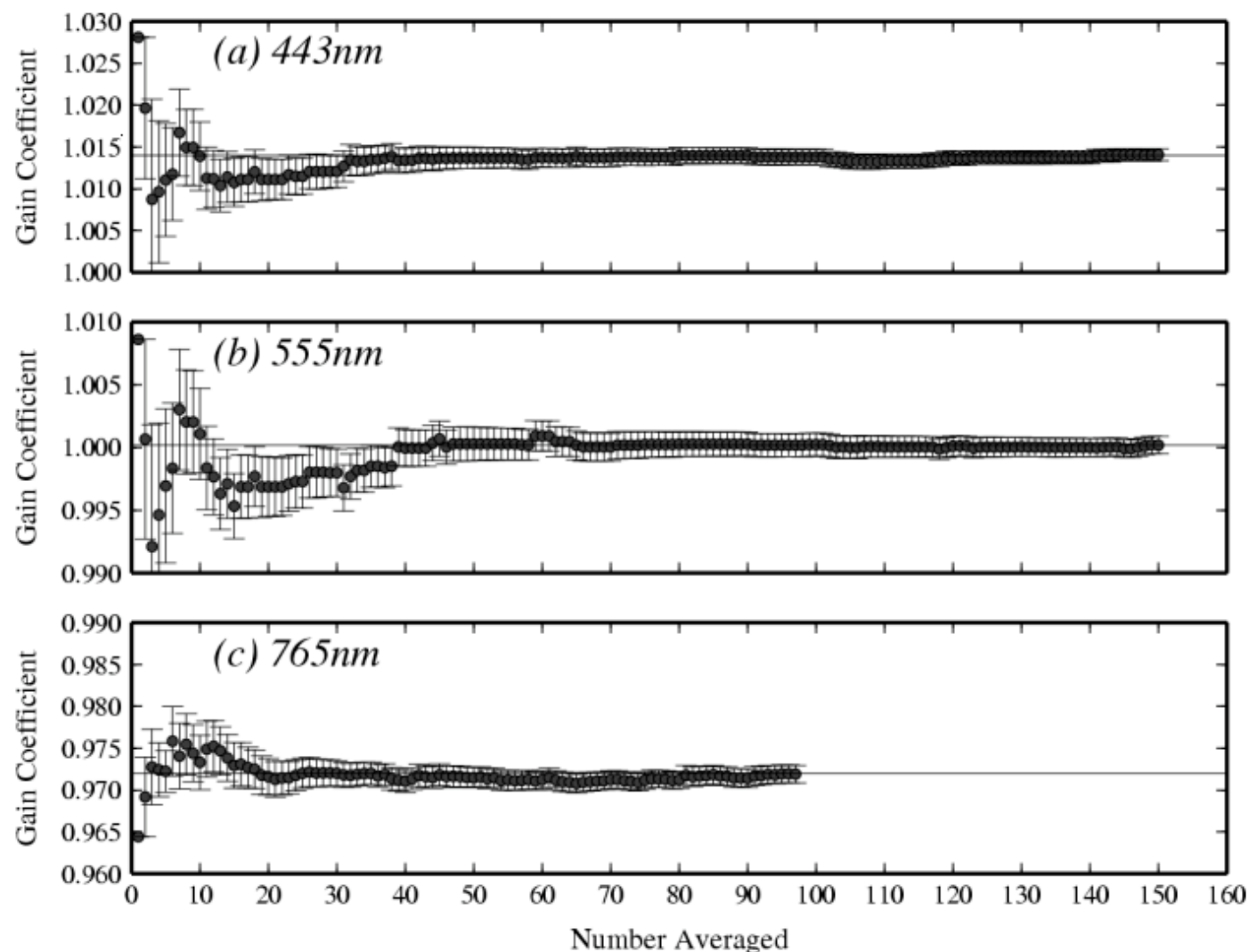


MOBY Water-leaving Radiance Time Series (1997-2000)



Vicarious Calibration Gain Convergence

- Only a small % of samples result in a MOBY-satellite “match-ups” for the vicarious calibration.
- For MODIS, took over 2 years to collect enough match-ups to derive gain corrections.



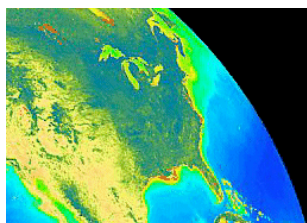
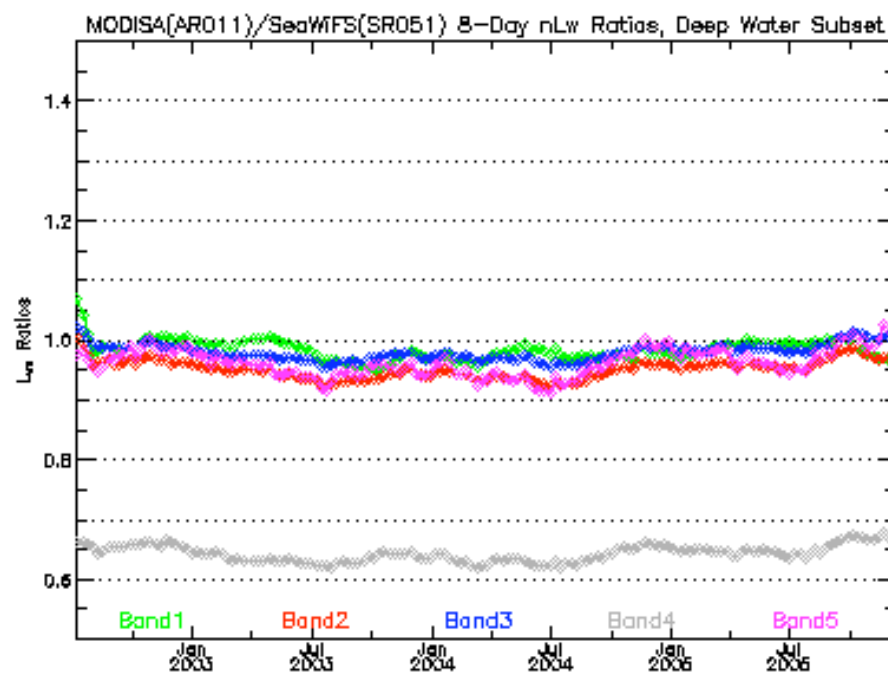
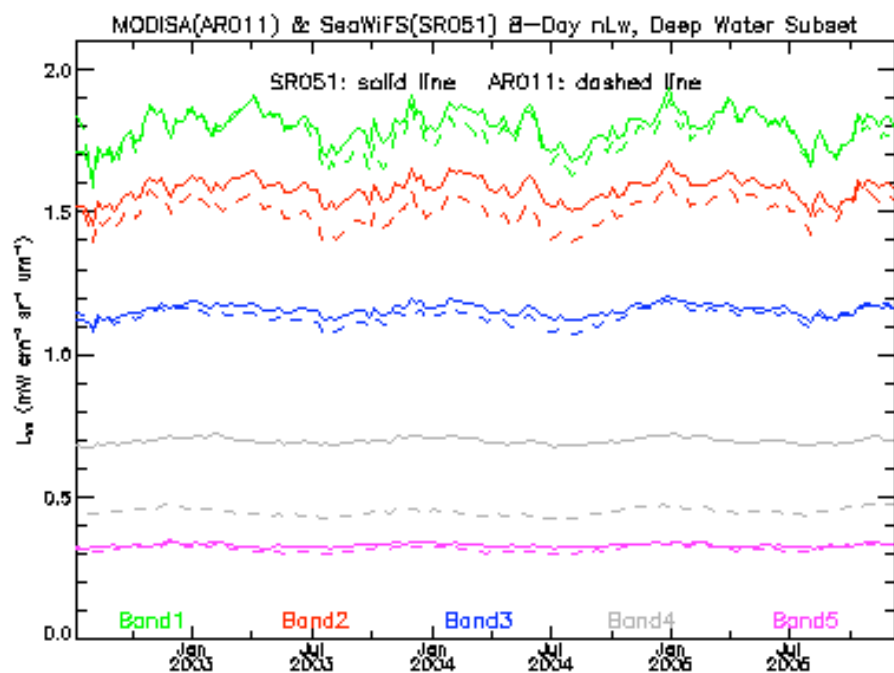
B. A. Franz, S. W. Bailey, P. J. Werdell, and C. R. McClain, "Sensor-independent approach to the vicarious calibration of satellite ocean color radiometry," *Appl. Opt.* 46, 5068-5082 (2007)

Direct Comparison of Satellite Lwn Retrievals

Global Deep-Water, 8-Day Composites, Common Bins

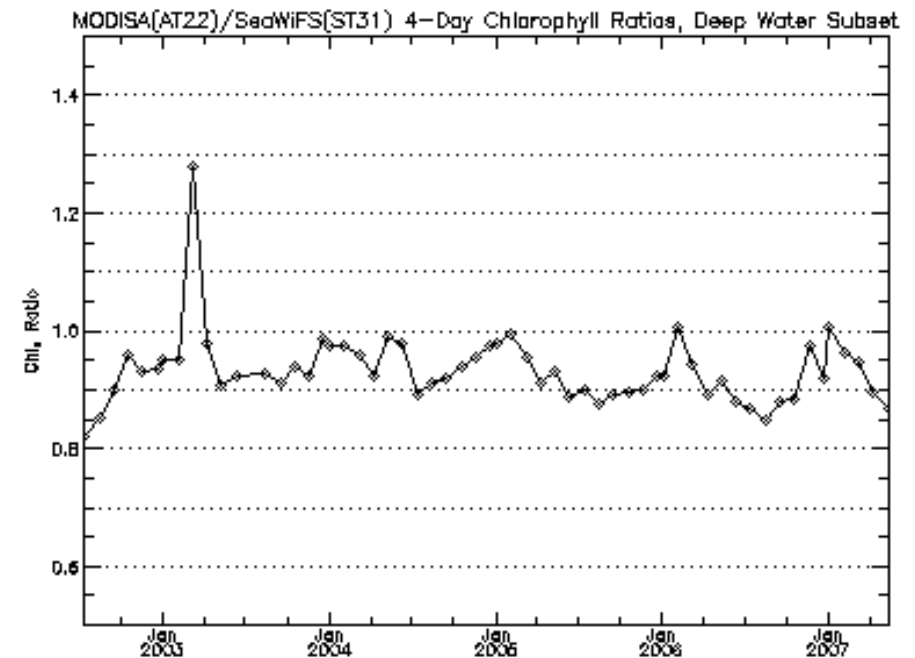
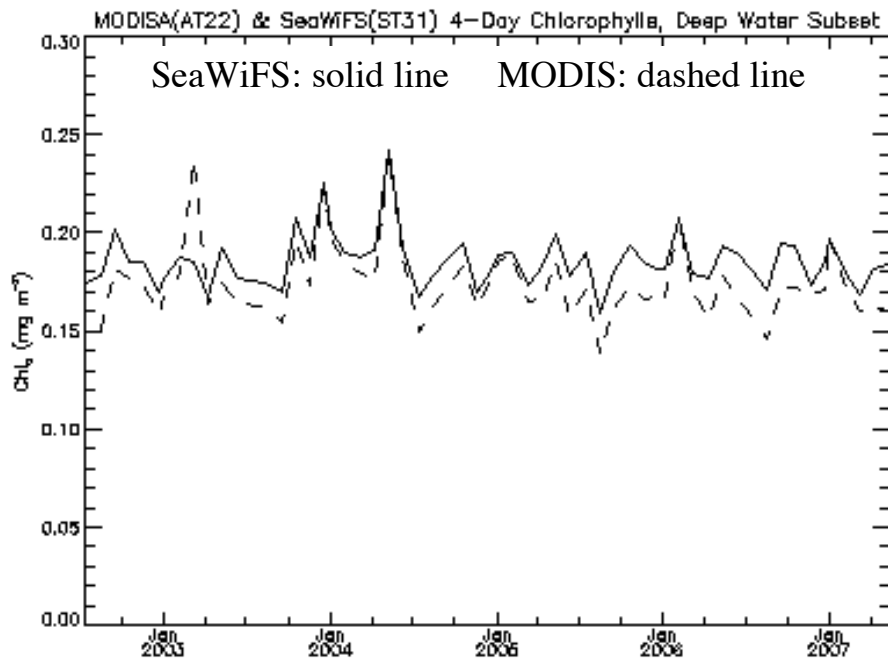
SeaWiFS & MODIS-Aqua

MODIS-Aqua / SeaWiFS

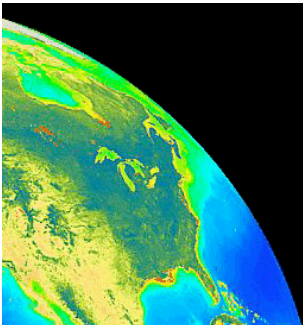


SeaWiFS-MODIS/Aqua Chl-a

Global Deep-Water* Means (2002-2007)

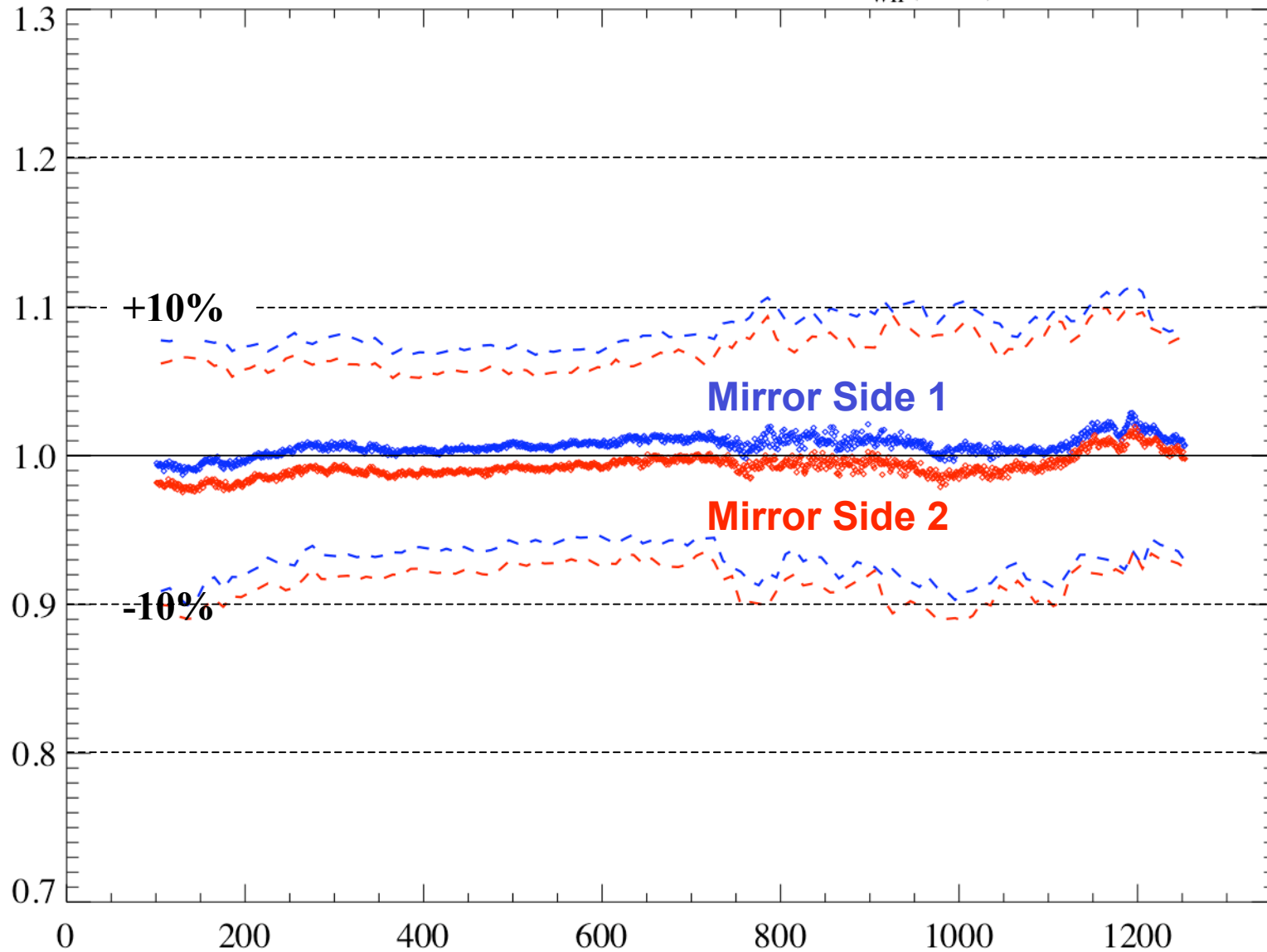


Ratio: MODIS-Aqua/SeaWiFS



MODIS-Aqua Residual RVS - Lwn(412)

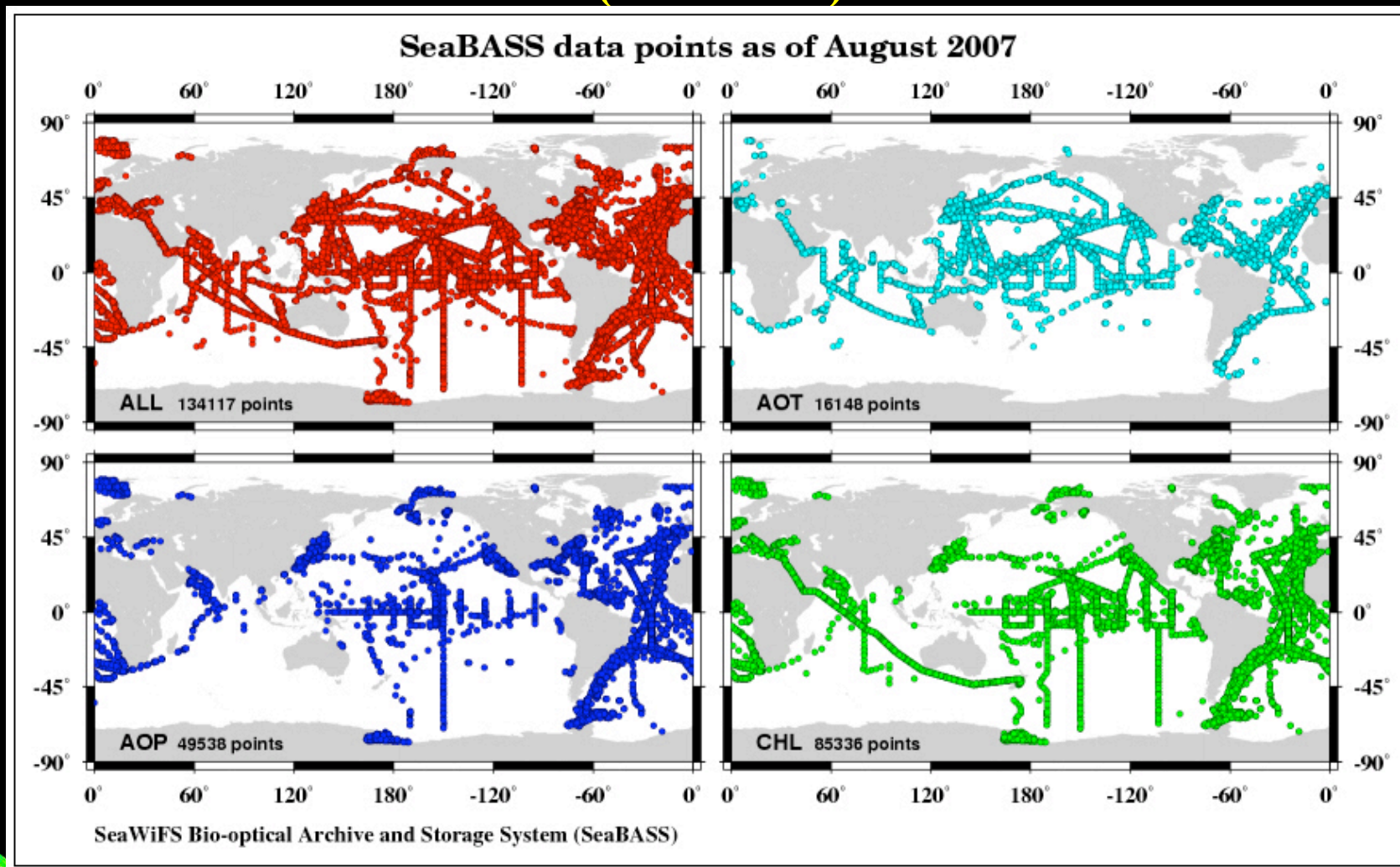
MODISA Global Mean Residual RVS in $L_{wn}(412)$ for 2006 266



$\frac{L2}{L3}$

Scan Pixel

SeaWiFS Bio-optical data Archive & Storage System (SeaBASS)



Data from over 1750 cruises

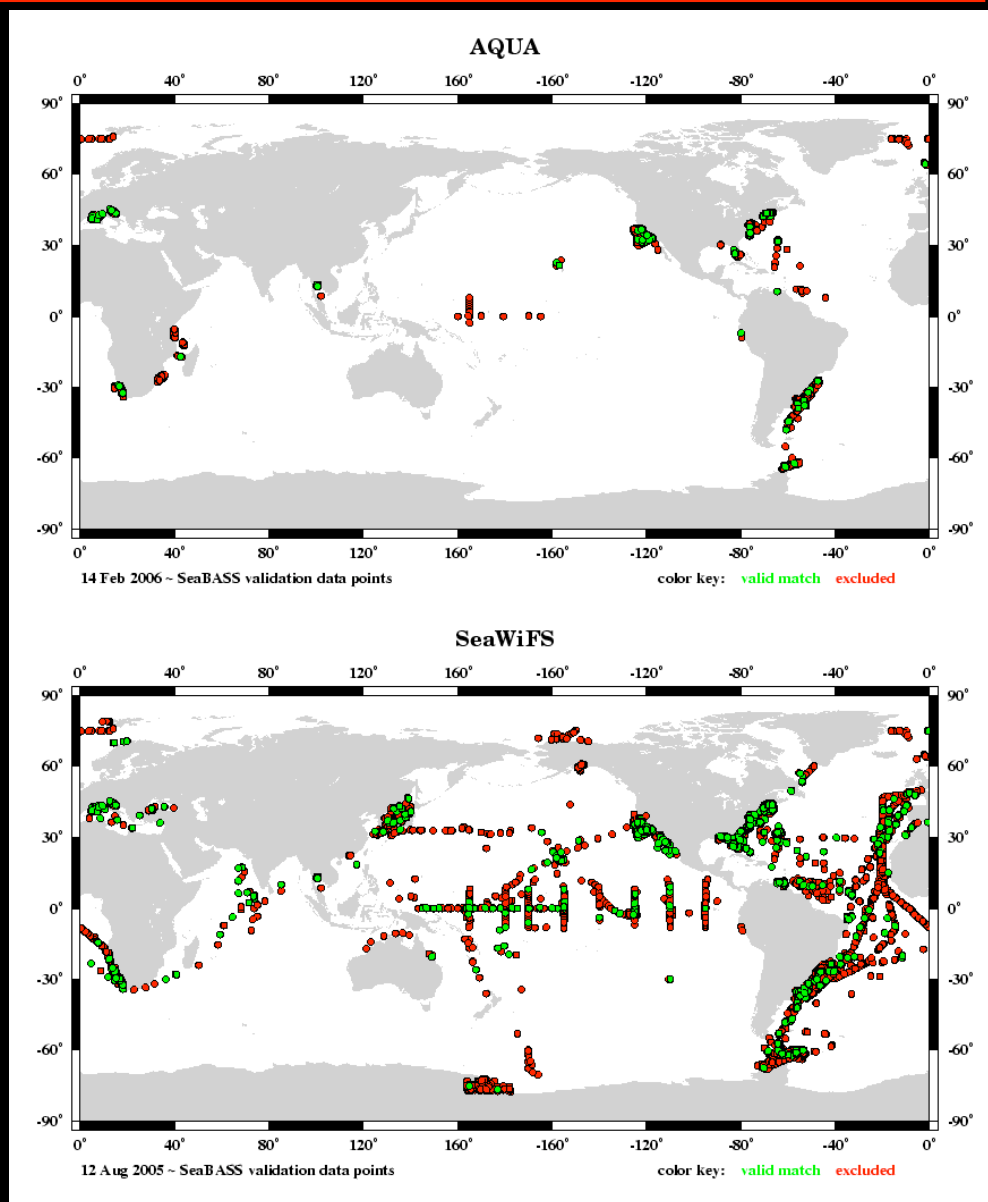
Apparent Optical Property (AOP); Chlorophyll-a (CHL); Aerosol Optical Thickness (AOT)

Available *In Situ* Match-Ups by Mission

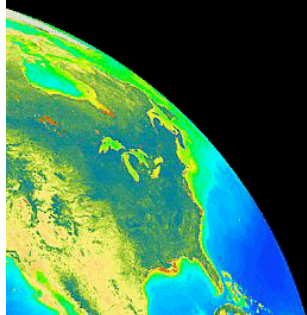
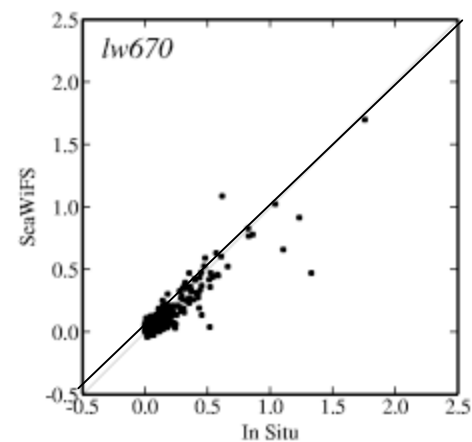
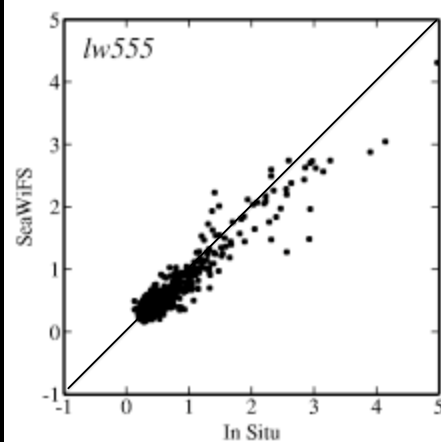
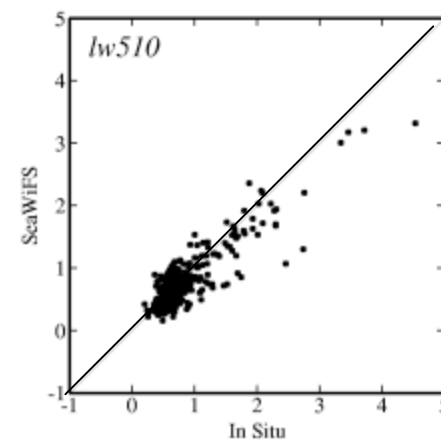
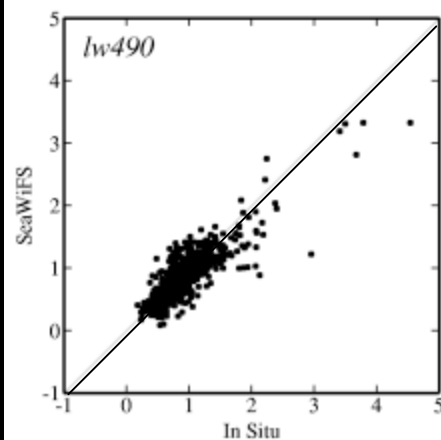
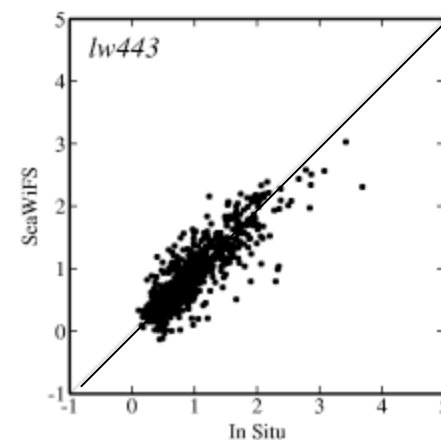
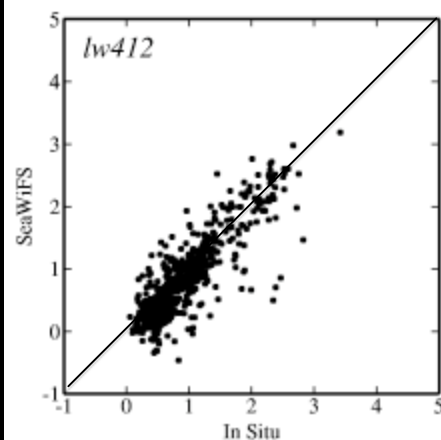
MODIS/Aqua
July 2002
- Present

Roughly 10% of stations
pass match-up criteria

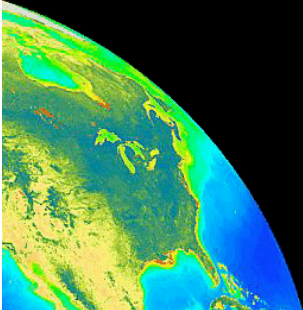
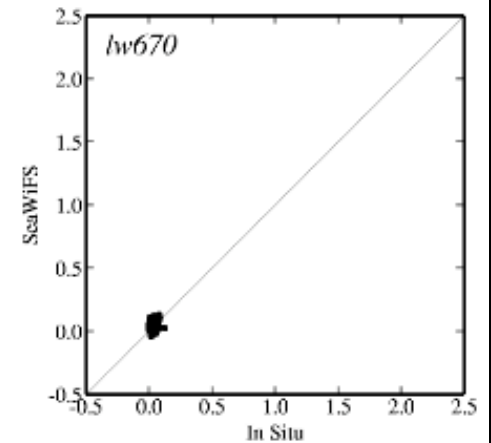
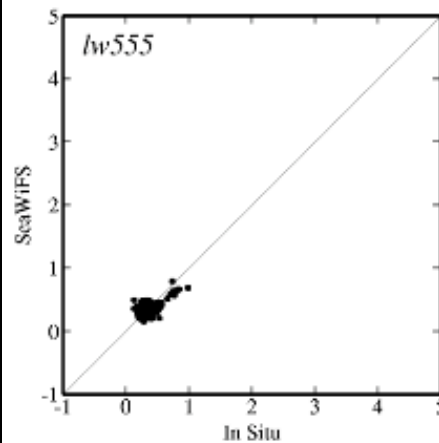
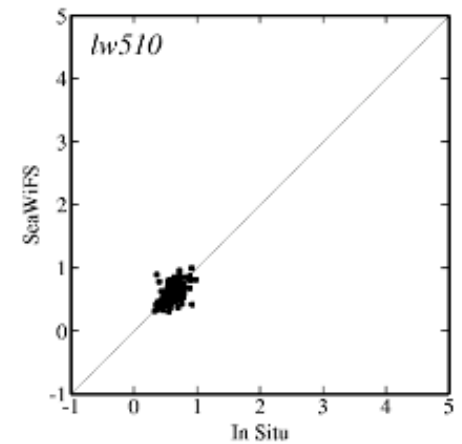
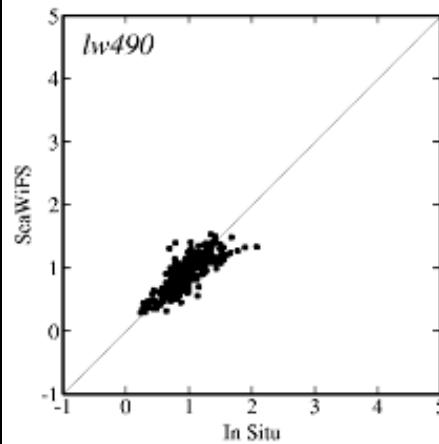
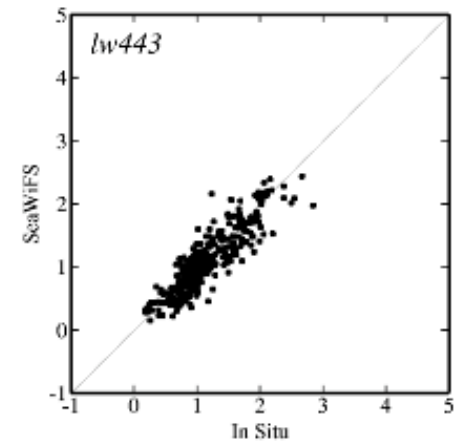
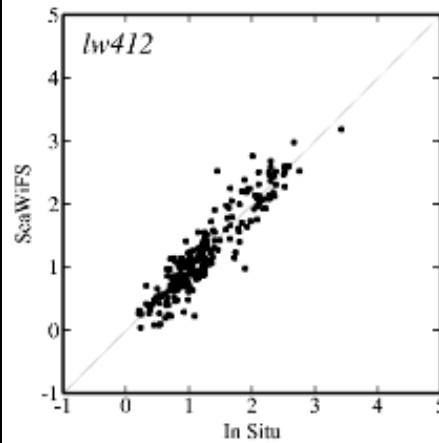
SeaWiFS
Sept 1997 -
Present



SeaWiFS Lwn Comparisons with Field Data

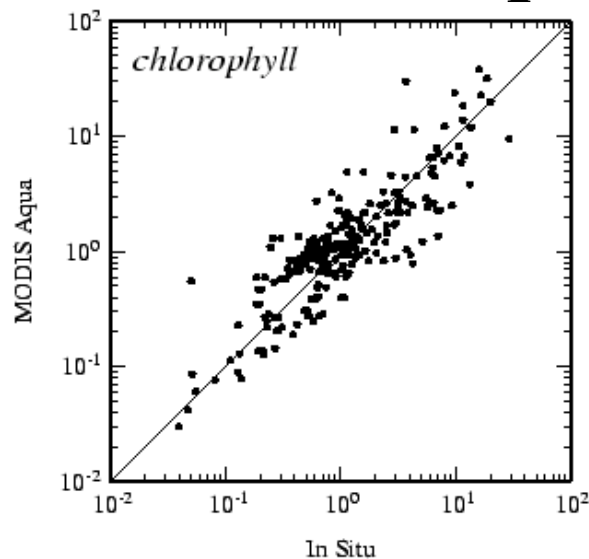


SeaWiFS Lwn Comparisons with Field Data: Deep Water Only (> 1000 m)

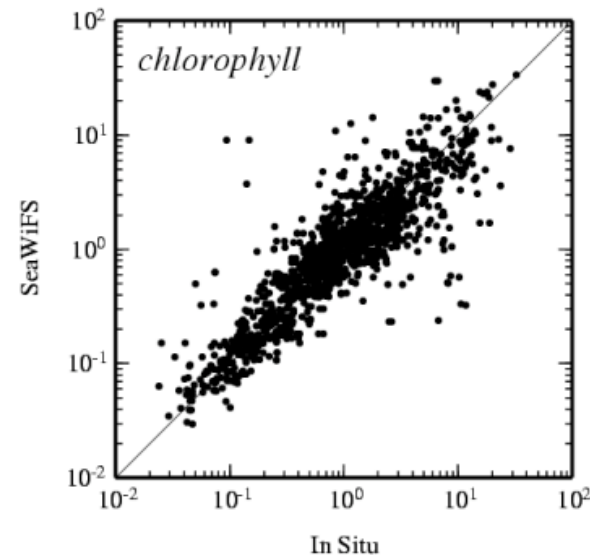


Comparison of Chlorophyll Retrievals to *In Situ*

MODIS/Aqua



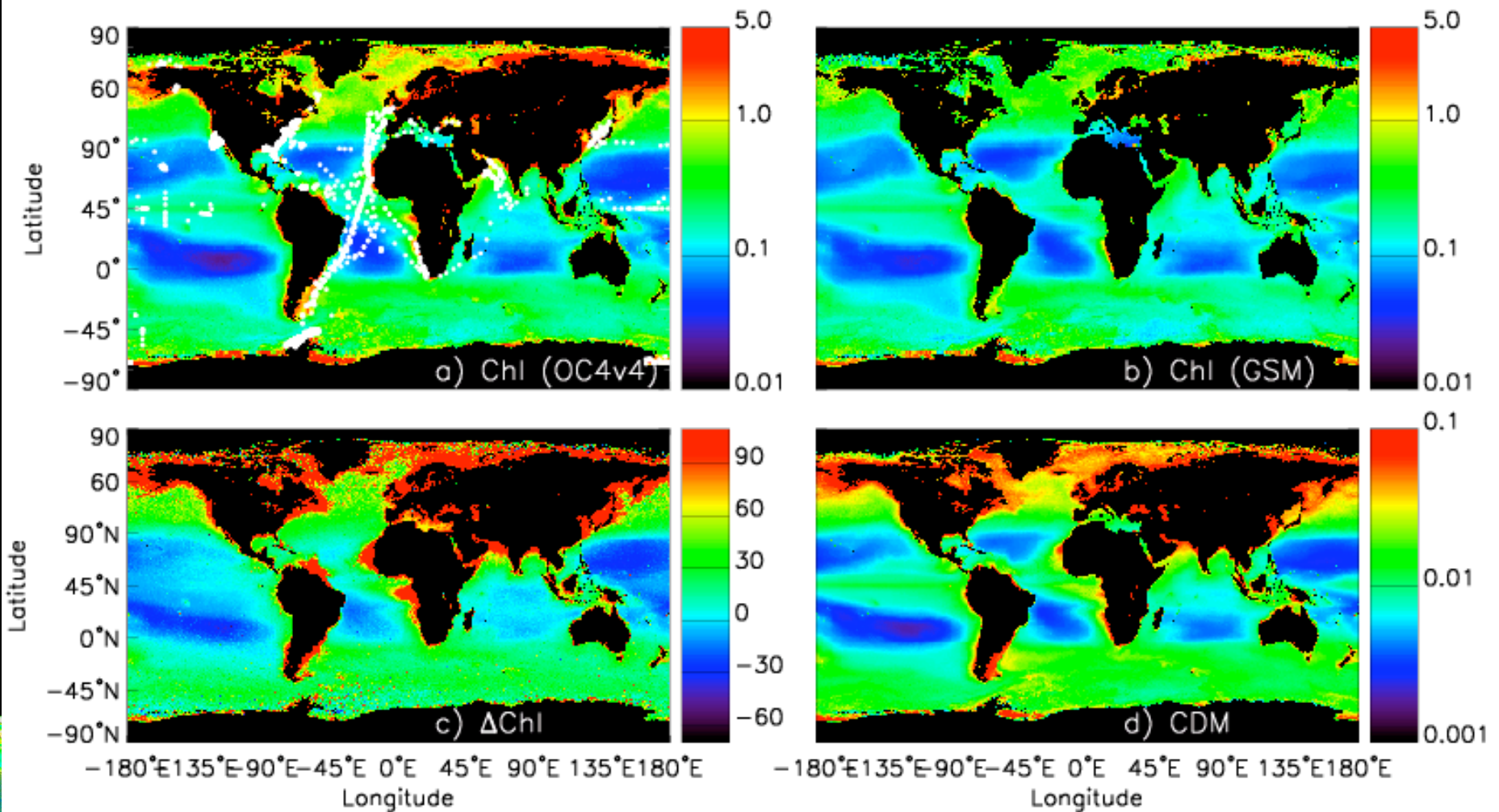
SeaWiFS



Sensor	# Matches	Mean Ratio	% Diff	r^2
SeaWiFS	1293	0.998	33.1	0.796
MODIS/Aqua	263	1.084	40.4	0.780

- Band-ratio algorithms very forgiving of radiometric biases.
- Semi-analytic reflectance inversion models very unforgiving
- ...required direction of future algorithm development.

Band Ratio (OC4) vs. Semi-analytic (GSM) Algorithms



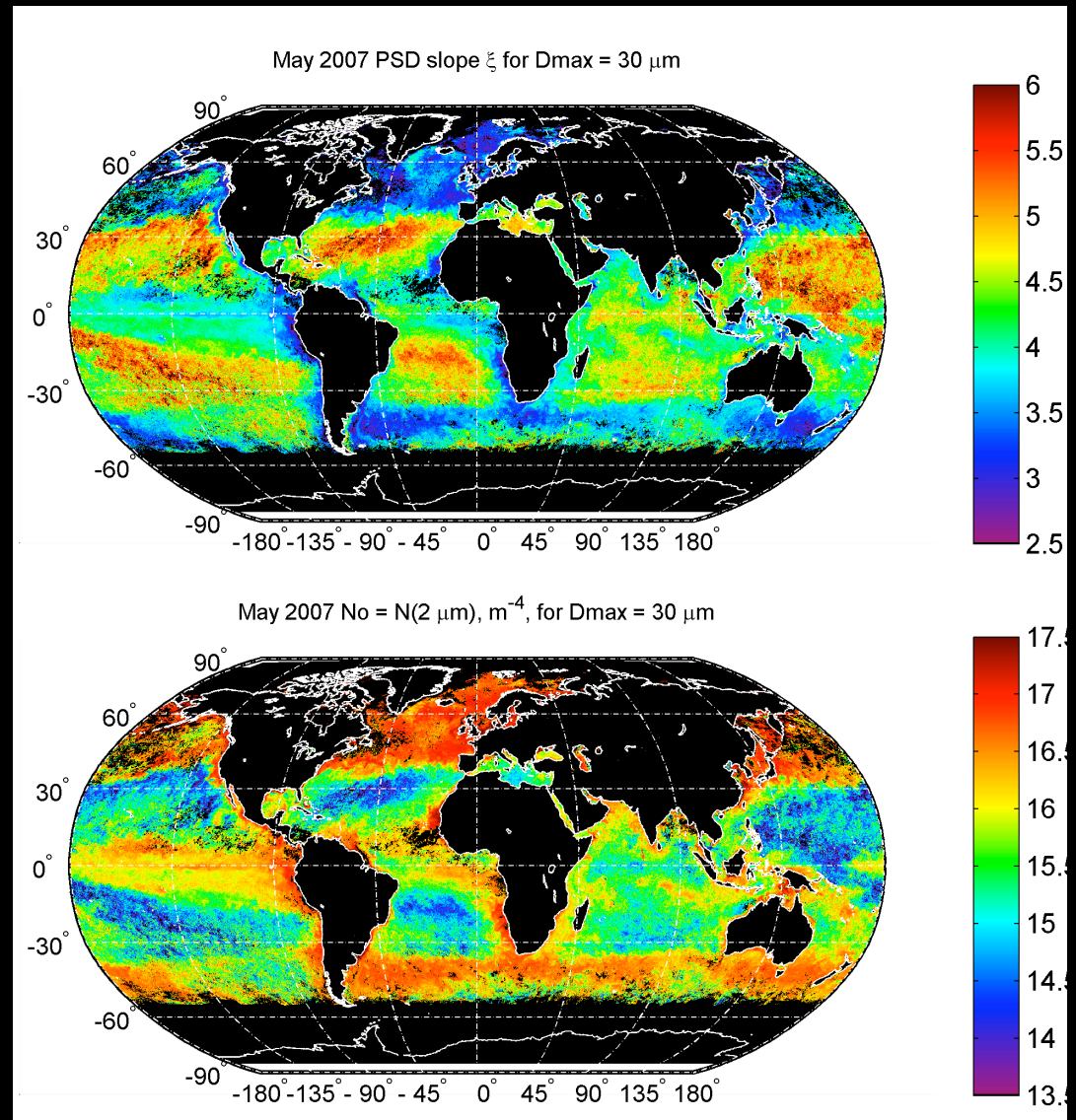
$$\Delta\text{Chl}(\%) = 100 * [\text{Chl}(\text{OC4v4}) - \text{Chl}(\text{GSM}) / \text{Chl}(\text{GSM})]$$

CDM: Colored Dissolved Matter

Semi-analytic Inversion Models: Marine Optical Properties

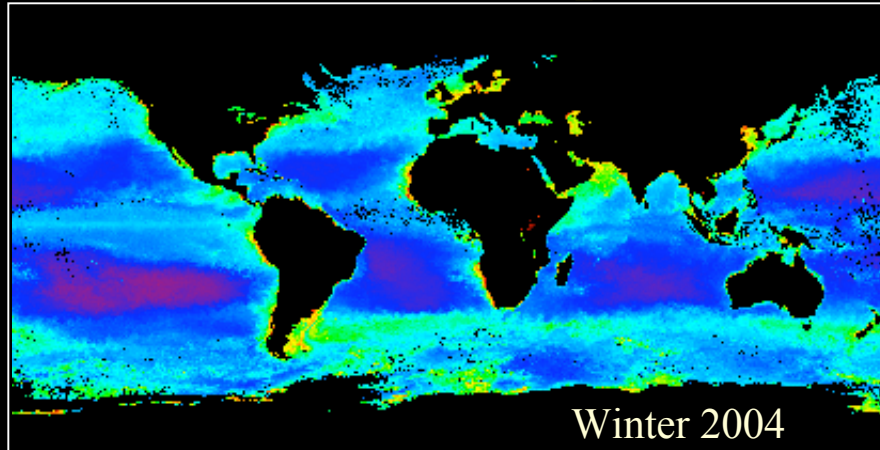
Marine
particle size
distribution
properties=>
phytoplankton
functional
groups, etc.

$\log_{10}(\text{particles}/\text{m}^{-4})$

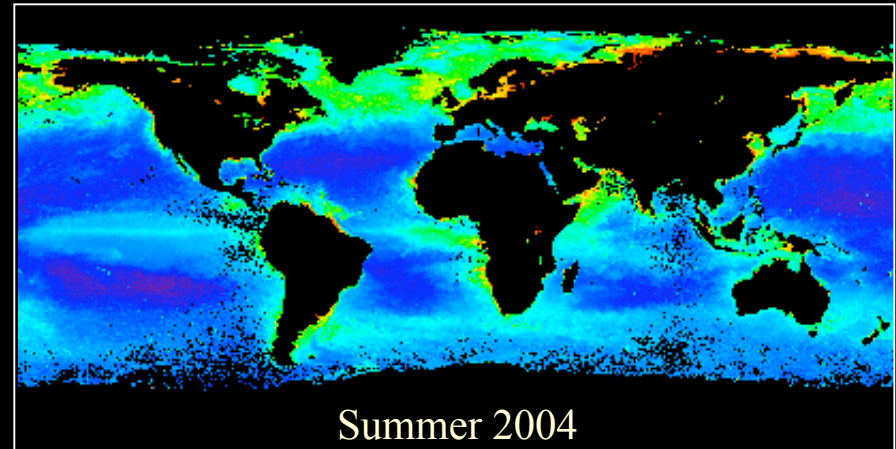
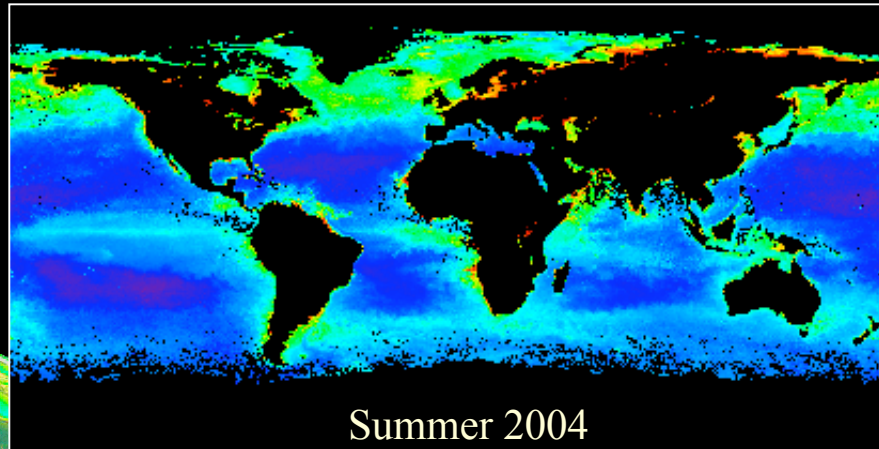
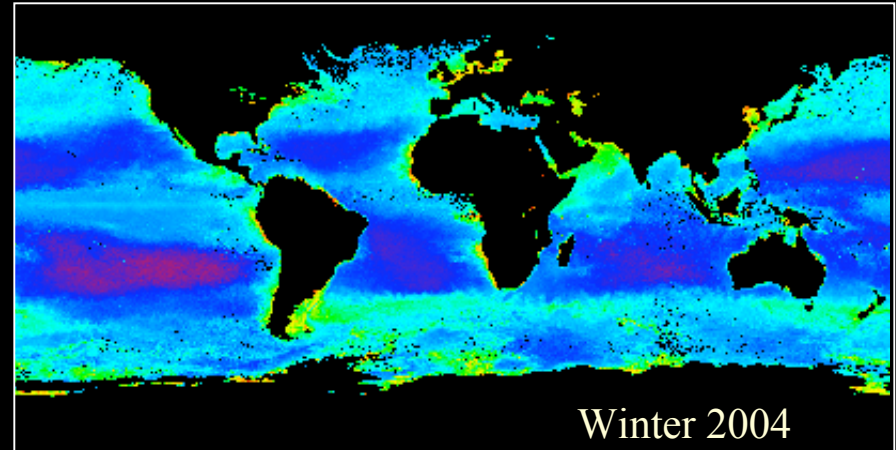


Seasonal Chlorophyll Images: Comparison across Sensors

MODIS/Aqua



SeaWiFS

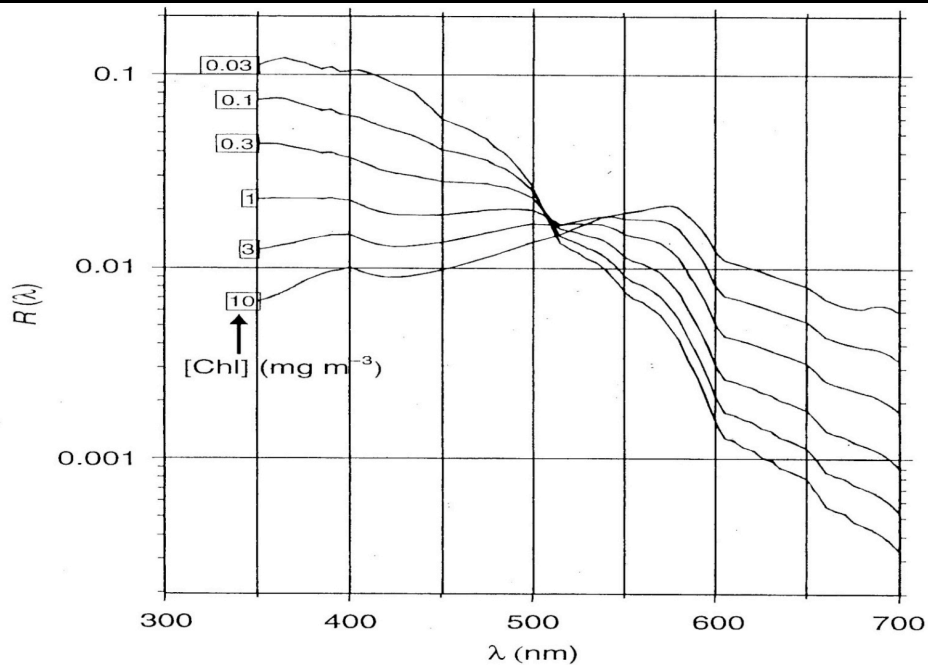


0.01-64 mg m⁻³

Ocean Reflectance: Open Ocean vs. Turbid Water

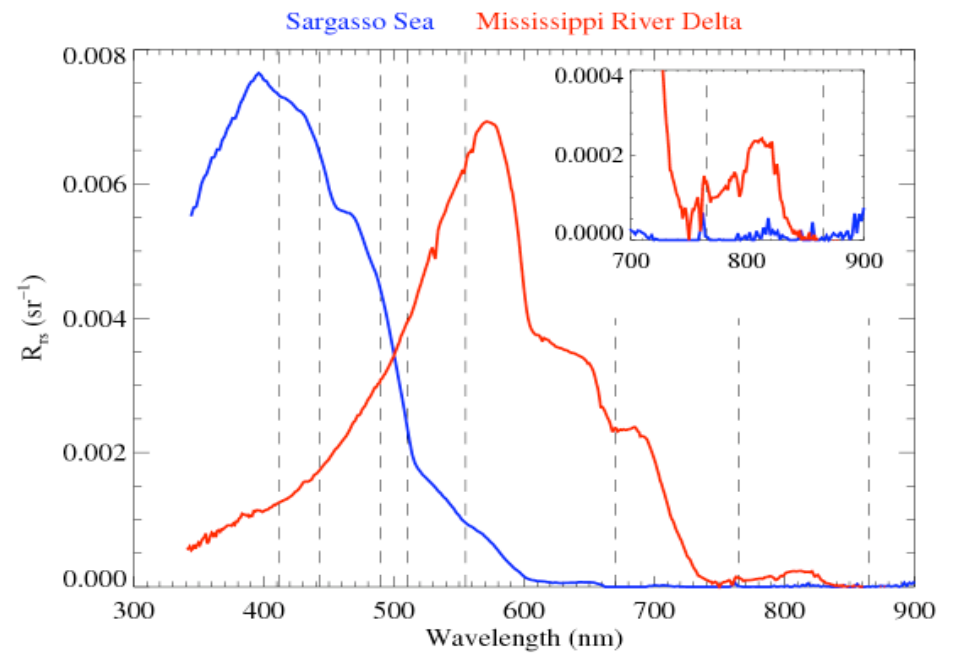
Irradiance Reflectance of the Ocean

$$R^0(\lambda) = E_u^{(0-)}(\lambda) / E_d^{(0-)}(\lambda)$$



Remote Sensing Reflectance $R_{rs} = L_w / E_d^{(0+)}$

Sargasso Sea and Mississippi River Delta



Atmospheric Correction Methodology (cont.)

- NIR Reflectance Correction in Turbid Water

-For higher concentration of chlorophyll (above 0.7 mg/m^3) the assumption that water-leaving radiance (L_w) in the NIR bands is zero is no longer valid. The correction is based on a bio-optical model that relates the remote sensing reflectance (R_{rs}) in the NIR as:

$$R_{rs}(\lambda) = R_{rs}(\lambda_0) [a_{tot}(\lambda_0)/a_{tot}(\lambda)] [b_b(\lambda)/b_b(\lambda_0)] -$$

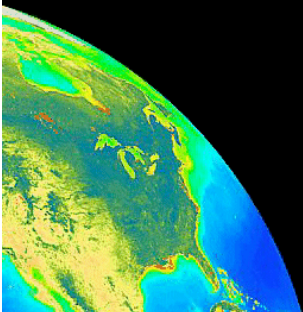
where $\lambda_0=670\text{-nm}$, and $\lambda=765$ and 865 nm

$$a_{tot}(\lambda) = a_w(\lambda) + a_{ph}(\lambda) + a_{dg}(\lambda) \quad \text{absorption coefficients}$$

$$b_b(\lambda) = b_{bw}(\lambda) + b_{bp}(\lambda) \quad \text{backscattering coefficients}$$

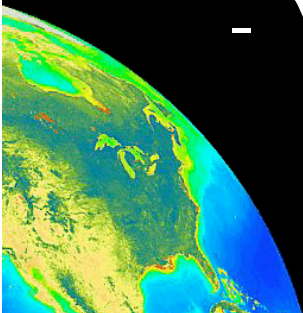
where $b_{bw}(\lambda) = a(S) \cdot \lambda^{-4.32}$ and $b_{bp}(\lambda) = m/\lambda + c$

$$R_{rs} = L_w / E_d, \text{ where } E \text{ is the surface solar irradiance.}$$



Atmospheric Correction Methodology (cont.)

- Model does not work well in high turbidity
- Alternative for turbid water aerosol corrections: **SWIR bands**
 - $L_w \sim 0$ even in most turbid waters due to extremely high water absorption
 - SeaWiFS: no SWIR bands
 - MODIS: 1240, 1640, 2130 nm bands
 - Aqua: 1640 band inoperative (bad detectors)
 - Low SNRs
 - VIIRS 1240, 1610 nm bands
 - SNRs only slightly higher than MODIS

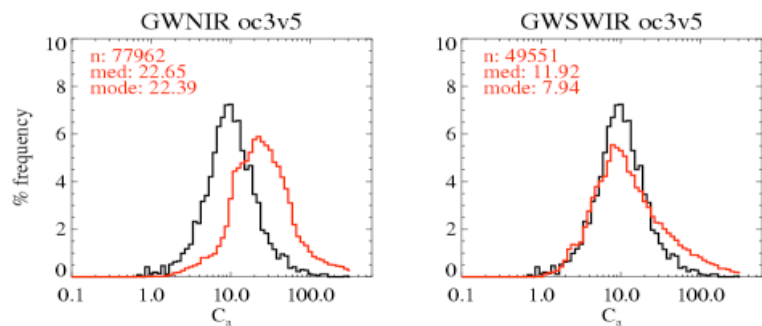


NIR

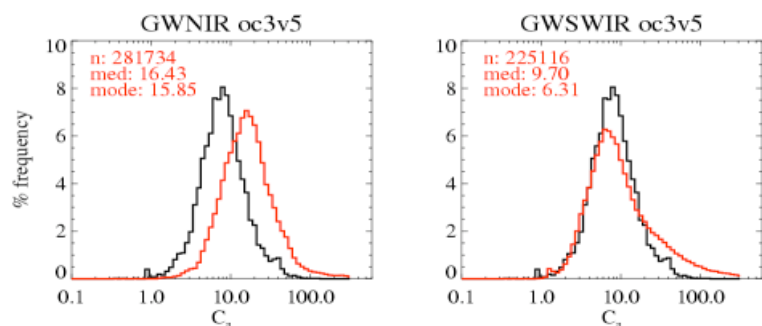
SWIR

Satellite vs In Situ

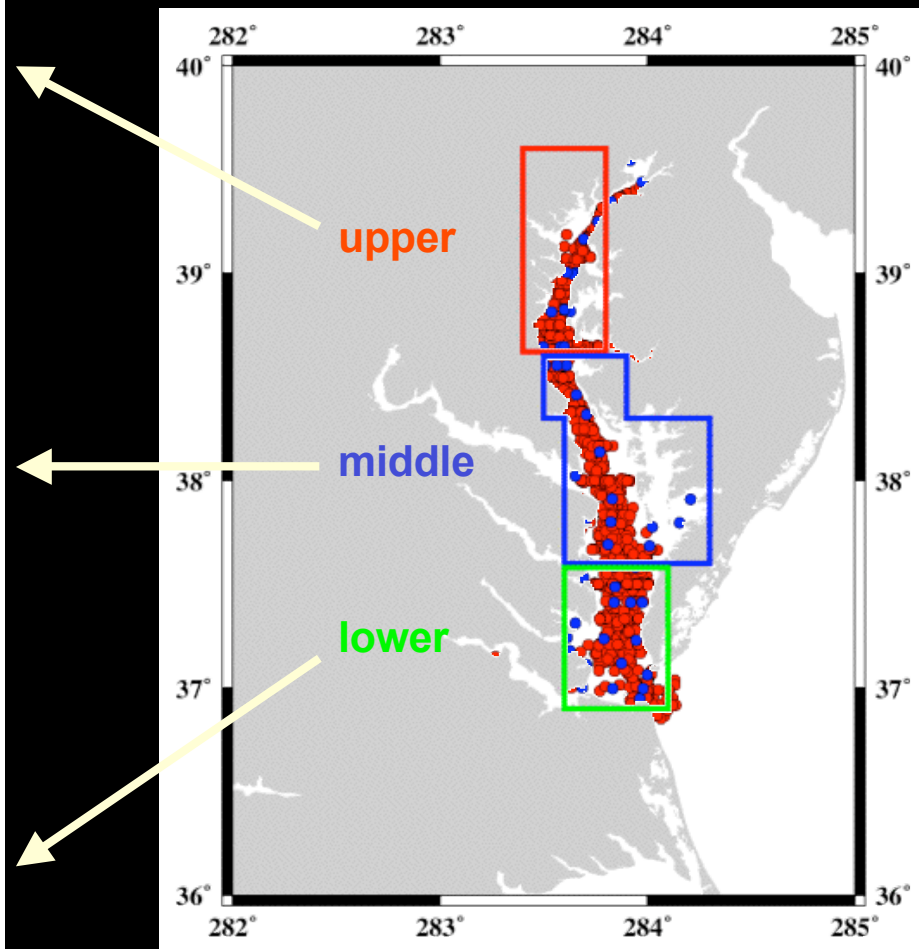
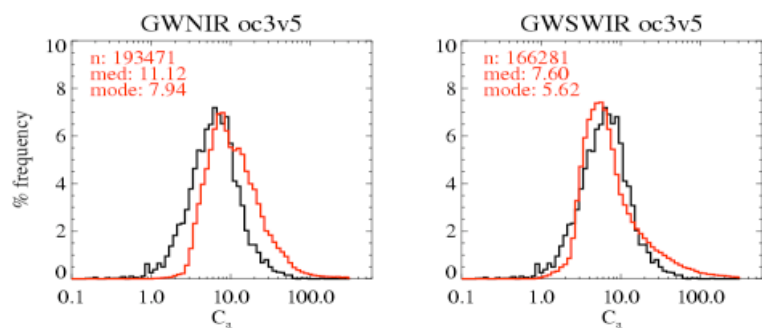
Upper Bay, ALL in situ = n: 3663, med: 10.52, mode: 10.00
color legend: in situ MODIS-Aqua



Mid Bay, ALL in situ = n: 5814, med: 8.43, mode: 7.94
color legend: in situ MODIS-Aqua



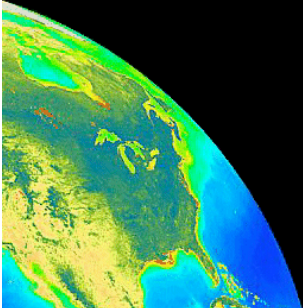
Lower Bay, ALL in situ = n: 7204, med: 6.50, mode: 6.31
color legend: in situ MODIS-Aqua



Chesapeake Bay

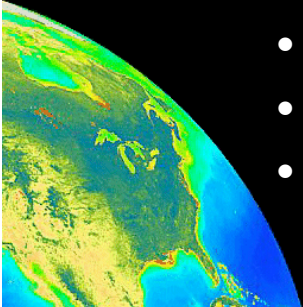
Atmospheric Correction Issues

- Absorbing aerosols
 - Concentrations below cloud mask reflectance threshold
 - No methodology for even flagging low concentrations
- NIR calibration
 - No reliable methodology for vicarious adjustment of 865 nm, 869 nm and MODIS SWIR bands
- Turbid water aerosol corrections
- NO₂ temporal & spatial coverage
 - For SeaWiFS, requires GOME, Sciamachy, and OMI cross-calibrated data.
 - Requires significant spatial interpolation to fill gaps



Future Directions

- Current sensors (SeaWiFS, MODIS, etc.)
 - Update aerosol models
 - New model suite being tested
 - Model properties consistent with coastal/island AERONET data
 - Evaluate utilization of Calipso & Glory data in SeaWiFS & MODIS atmospheric corrections, especially for absorbing aerosols
- ACE
 - Absorbing aerosol identification & atmospheric corrections
 - Combination of lidar and polarimeter aerosol height & type inputs
 - UV band aerosol corrections
 - Use of UV bands for coastal/turbid water aerosol corrections
 - Assumes insignificant reflectance
 - Coupled ocean-atmosphere RT-based aerosol corrections
 - Several models published to date.
 - Computationally too demanding for routine global processing at this time
 - Performance improvement not dramatic except in presence of absorbing aerosols

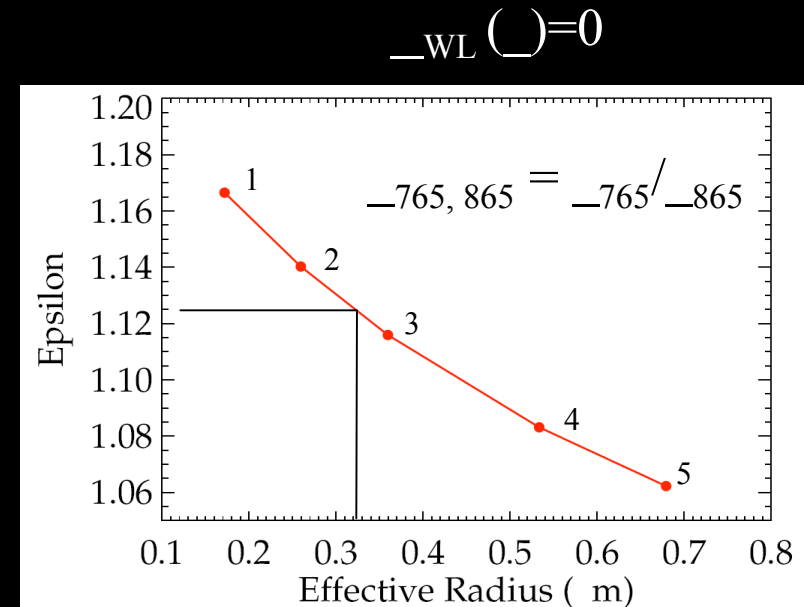
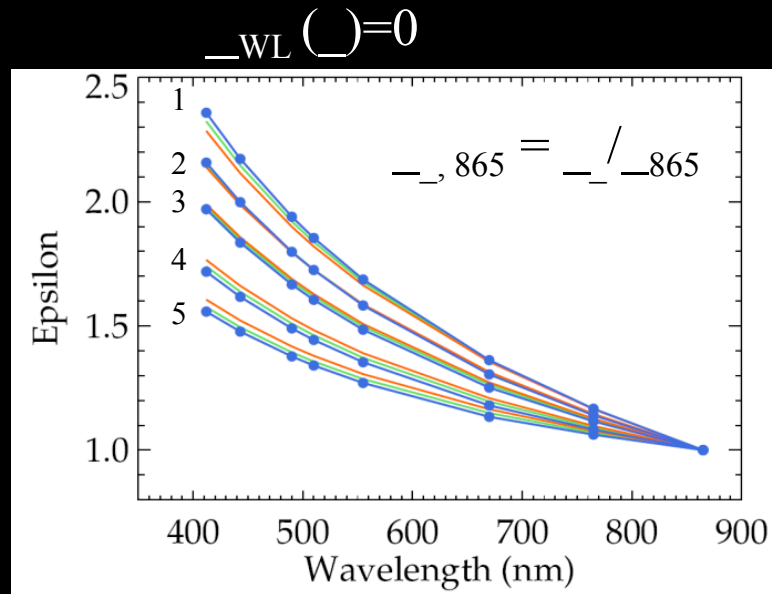


Atmospheric Correction Methodology

• Aerosol Model Selection

- OC processing uses NIR measurements to select aerosol model

$$\tau_{765, 865} = \tau_{765} / \tau_{865}$$

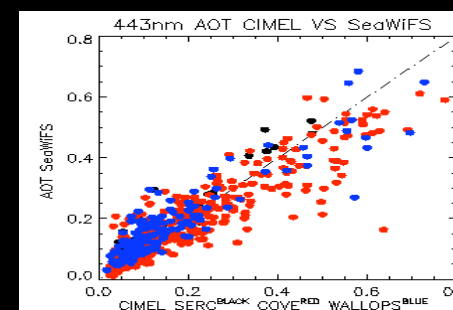
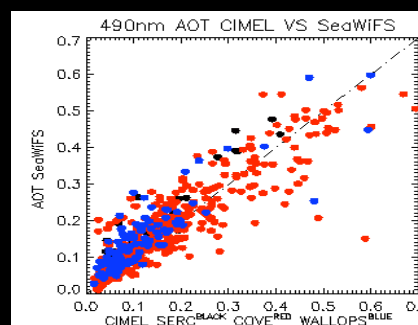
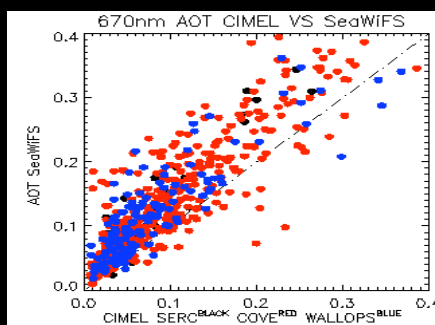
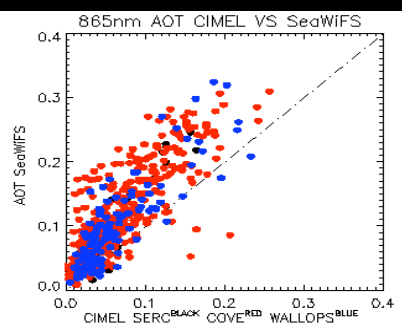


- Select two models that bracket the observed $\tau_{765, 865}$
- In operational processing we use ratio of single-scattering-reflectance values to compute $\tau_{765, 865}$

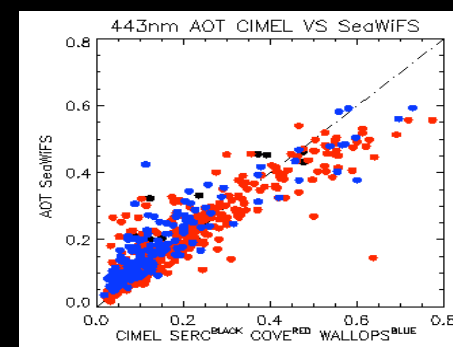
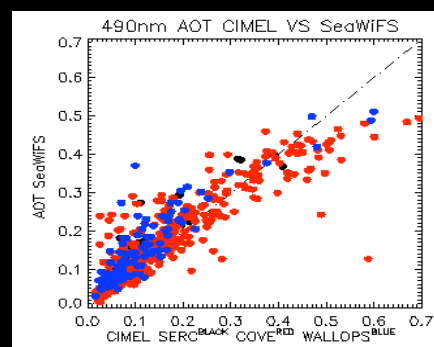
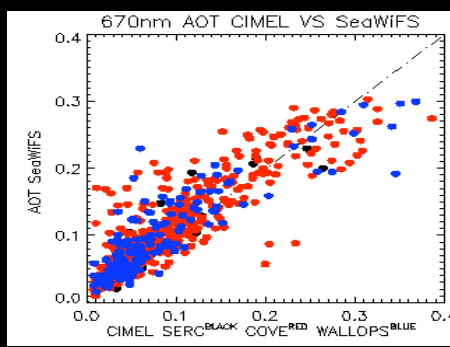
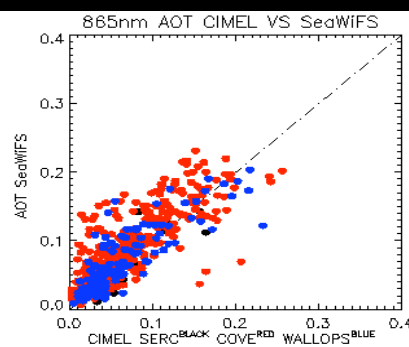
Aerosol Opt. Thickness (MODIS vs. AERONET)

Comparisons from 3 AERONET sites around Chesapeake Bay

AOT Based on Operational Models



AOT Based on New Models



865 nm

670 nm

490 nm

443 nm

AOT range: 0 – 0.7

BACKUP SLIDES



Aerosol Optical Thickness Retrievals

Similar bias found
in MODIS/Aqua
retrievals.

